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Riverine flux of dissolved phosphorus to the coastal sea may be overestimated, especially in estuaries of gated rivers: Implications of phosphorus adsorption/desorption on suspended sediments

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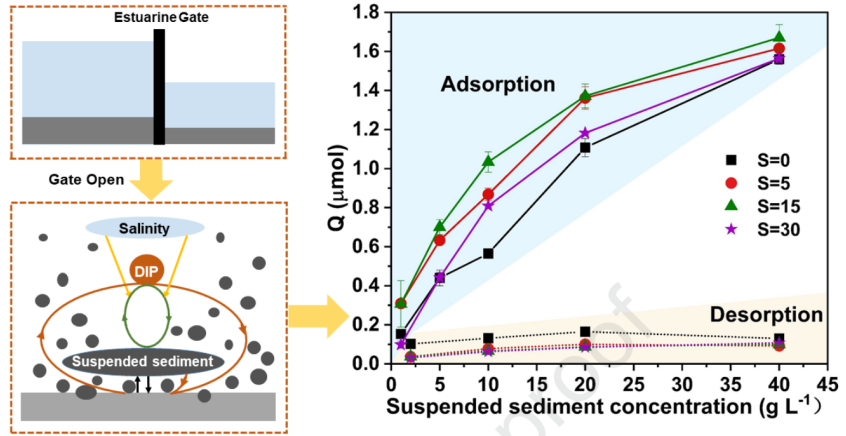
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1 Abstract

2 The flux of terrestrial dissolved inorganic phosphorous (DIP, i.e. PO_4^{3-}) via rivers
3 into coastal seas is usually calculated by simply multiplying its concentration with the
4 corresponding water flow at the river mouth. Subsequent adsorption/desorption of
5 DIP onto suspended sediment and the influence of salinity in the estuary are often
6 overlooked. A series of DIP adsorption/desorption experiments under different
7 salinities (0, 5, 15, 30) and suspended sediment concentrations ($1\text{-}40 \text{ g L}^{-1}$) were
8 conducted in order to assess the potential influence of these factors on the overall DIP
9 loading to the coastal zone. The effect of different sea-salt ions on DIP
10 adsorption/desorption was also assessed by comparing different experimental
11 solutions (NaCl solution, artificial seawater and real seawater). In estuaries, the
12 adsorption of DIP to suspended sediments was greater than desorption, and the net
13 adsorption increased with increasing concentration of suspended sediments and
14 salinity. This enhanced DIP adsorption onto suspended sediment reduces the riverine
15 discharge of DIP to coastal ecosystems. Disregarding this process, especially for the
16 gated estuaries with high sediment resuspension, potentially leads to an
17 overestimation of the terrestrial DIP input to the coastal region.

18 **Key words:** seaward dissolved inorganic P flux; adsorption/desorption;
19 suspended sediment; salinity

20

21

22 1. Introduction

23 Anthropogenic pressure on estuarine and coastal environments is increasing,
24 due to the growing population along the coast (CIESIN, 2018). Coastal eutrophication
25 is one of the most severe and wide-spread environmental problems, primarily caused
26 by intensified influx of riverine nutrients (phosphorus and nitrogen) (Ballagh et al.,
27 2019). Hence, quantification of nutrient inputs from land to sea is important for
28 understanding and managing coastal eutrophication.

29 Typically, the nutrient flux to the sea is calculated by multiplying nutrient
30 concentrations at the river mouth with corresponding water flow. However, this
31 approach neglects the important physicochemical processes taking place in the
32 estuarine systems, such as adsorption and desorption of nutrients on suspended
33 sediment. These processes are highly affected by changing physicochemical
34 conditions along the estuarine-coastal gradient. The calculated nutrient flux may
35 therefore not represent an “effective” input to the coastal zone. These processes must
36 therefore be taken into account when estimating the riverine loading of a pollutant,
37 especially regarding particle-reactive chemical species such as phosphate (PO_4^{3-}).
38 This is particularly the case for estuaries with river inflow controlled by gates,
39 signified by pulsed flow patterns and high sediment resuspension. The complex
40 chemical and hydrodynamic environment in estuaries fed by gated rivers affects the P
41 cycling processes and thereby the DIP flux to the sea (Wang et al., 2010; Zhang et al.,
42 2019a; Meng et al., 2014a). Although the total influx from gated rivers is lower than
43 that from rivers without gates or dams, the increased hydrodynamic turbulence during
44 discharge pulses enhances sediment resuspension, especially during the flood season
45 (Duan et al., 2008; Lehner et al., 2011; Zarfl et al., 2015; Eiriksdottir et al., 2017;
46 Ding et al., 2019). According to field investigation and other studies, the highest
47 measured sediment concentration in gated rivers or the maximum estuarine turbidity
48 zones is around 40 g L⁻¹ (Uncles et al., 2006; Snedden et al., 2007; He et al., 2010;
49 Yao et al., 2016; Palinkas et al., 2019).

50 Phosphorus (P) is an essential element for phytoplankton growth and DIP
51 typically limits primary production during spring in estuaries receiving high inputs of
52 nitrogen (Ruttenberg, 1992; Ballagh et al., 2019). Rivers are the dominant source of P
53 to coastal seas (Benitez-Nelson, 2000). Accurate estimates of the riverine flux of
54 bioavailable P (i.e., mainly dissolved inorganic phosphate, DIP) is a prerequisite for
55 selecting appropriate abatement strategies to ensure a sustainable management of the
56 coastal environment (Ding et al., 2019; Palinkas et al., 2019; Liu et al., 2020).

57 The transport and transformation of P through an estuary is governed by a
58 complex set of processes. In addition to a large number of environmental factors in
59 the water column affecting DIP concentrations, processes at the sediment-water
60 interface also influence P transformations and thereby the P flux (Wang et al., 2012;
61 Kang et al., 2018; Rapin et al., 2019). For example, oscillations in organic matter
62 content and redox enhance P release from sediment (Wang et al., 2012; Kang et al.,
63 2018), while submerged macrophytes will enhance DIP uptake (Wang et al., 2012). In
64 estuaries, the interaction between DIP in water and suspended sediments is a key
65 process controlling bioavailable P. Suspended sediments can efficiently scavenge DIP
66 from the water column through adsorption, thereby decreasing the DIP flux to the sea
67 (Millero et al., 1986; Millero et al., 2001; Liu et al., 2002; Egger et al., 2015). On the
68 other hand, DIP released from the sediment via desorption may subsequently increase
69 the DIP flux to the sea. For gated estuaries, it remains unclear whether enhanced
70 sediment resuspension actually increases or decreases the net seaward DIP flux.

71 Adsorption of DIP onto suspended sediment is influenced by salinity,
72 temperature and pH, as well as the characteristics of the sediments such as texture and
73 organic matter content (Jones et al., 1993; Sundareshwar et al., 1999; Cao et al., 2011;
74 Zhang et al., 2011; Zhou et al., 2018). Among these factors, salinity is the most
75 prevalent factor affecting the adsorption process in estuaries, due to the large
76 gradients in salinity from fresh to seawater. The estuarine gradients in temperature,
77 organic matter and pH are generally smaller and to a large extent controlled by

78 salinity (Jun et al., 2013; Flower et al., 2017; Paudel et al., 2019). The fact that
79 salinity enhances phosphate adsorption implies that the phosphate oxyanion is bound
80 to active sites via inner-sphere (ligand exchange) complexation. Electrostatic outer-
81 sphere interactions between phosphate and suspended sediments are weakened by
82 increased salinity due to competition for active sites by electrolyte anions, such as
83 chloride (Bai et al., 2017; Huang et al., 2015). Moreover, the inner-sphere
84 complexation is facilitated by the high sodium ion (Na^+) concentration weakening the
85 electrostatic outer-sphere interactions through lowering of particle net negative
86 surface charge (Zhang et al., 2019b). On the other hand, increased salinity also
87 promotes particle flocculation (Asmala et al., 2014; Bai et al., 2017), which decreases
88 the total surface area of the suspended particles and thereby their adsorption capacity.

89 Solutions of sodium chloride (NaCl) are commonly used instead of real
90 seawater to analyze salinity effects on DIP adsorption (Wang et al., 2006; Wu et al.,
91 2012; Bai et al., 2017; Zhang et al., 2019b; also see Fig. S1). In addition to sodium
92 (Na^+) and chloride (Cl^-), seawater contains significant concentrations of calcium
93 (Ca^{2+}), magnesium (Mg^{2+}), sulfate (SO_4^{2-}), and bicarbonate (HCO_3^-) (Riebesell et al.,
94 2010). The influence of these additional ions on the DIP adsorption differs from that
95 of Na^+ and Cl^- (Flower et al., 2017). Therefore, it is of special importance to assess
96 the specific ionic effect on the estimates of the DIP seaward flux. This is especially a
97 concern in estuaries where the load of suspended sediments is high, such as from
98 rivers with gates.

99 This paper aims to evaluate the influence of suspended sediments on the seaward
100 DIP flux, based on experiments for assessing the effect of salinity, ionic composition
101 and sediment concentration on DIP adsorption and desorption processes, especially
102 the scavenging effect caused by the elevated sediment resuspension in gated rivers.

103

104 **2. Methods**

105 **2.1. Sampling**

106 An estuarine surface sediment (ca. 0-5 cm) bulk sample was collected using a
107 plastic shovel in November 2019 from the DuLiuJianHe River, Tianjin, China, above
108 a tidal gate (38°45'58.25" N; 117°33'57.76" E) close to the river mouth. The
109 sediment sample was air-dried at room temperature (~ 27 °C) to a constant weight,
110 then softly ground to crush aggregates and sieved through 30 mesh.

111 A real seawater sample, with salinity of 30, was collected with a plastic bucket in
112 January 2020 from Dongjiang Port (39°1'8.38" N; 117°49'5.80" E), Tianjin, China.
113 In the laboratory, the seawater was filtered through 0.45 µm filter to remove
114 suspended particles, and stored in clean polyethylene bottles at 4 °C for later
115 experiments.

116 Basic physical and chemical characteristics of the collected sediment and
117 seawater samples are presented in Table S1.

118 **2.2. Experiments**

119 A solution with the suspended sediment sample in real seawater and NaCl
120 solution were used to conduct the P adsorption and desorption experiments. To
121 simulate the change in salinity from freshwater to seawater, solutions with a salinity
122 gradient of 0, 5, 15 and 30 were prepared by diluting the real seawater, NaCl solution
123 and artificial seawater with deionized water. For simplicity, the aqueous solutions
124 made from real seawater, NaCl solution and artificial seawater are referred to as
125 seawater, NaCl solution and artificial seawater, respectively. Four laboratory
126 experiments were conducted: 1) Adsorption; 2) Desorption; 3) Kinetics; and 4) Salt-
127 specific effects. All mixtures were shaken at 130 rpm for 48 h in a constant
128 temperature-water shaker at 27 °C, to achieve adsorption equilibrium. The suspensions
129 were centrifuged at 6000 rpm for 10 min, and the supernatants were filtered through
130 0.45 µm filter. The DIP content in the filtrates were measured spectrophotometrically
131 using the ammonium-molybdate method (Hansen and Koroleff, 2007; Murphy and
132 Riley, 1962).

133 **1) Adsorption experiments**

134 For the phosphate adsorption on suspended sediment, 39 mL solutions with
135 different salinities were added 1 mL 2 mM KH_2PO_4 solution, generating a P
136 concentration of 50 μM . A maximum sediment concentration of 40 g L^{-1} was used in
137 this study. A series of adsorption experiments were conducted using solutions (50 μM
138 P) with sediment levels of 1, 5, 10, 20 and 40 g L^{-1} under different salinities (0, 5, 15,
139 30). Each treatment was measured in duplicate.

140 **2) Desorption experiments**

141 No extra P was added to the solutions. Besides from this, the DIP desorption
142 experiments were conducted in a similar manner as for the adsorption experiments
143 with sediment concentrations at 2, 10, 20, and 40 g L^{-1} and salinities at 0, 5, 15 and 30.

144 **3) Kinetic experiments**

145 The DIP concentration was reduced to 20 μM in the kinetic experiments in order
146 to study the DIP adsorption efficiency on sediment devoid of any adsorption capacity
147 limitations. Salinities and concentrations of suspended sediment tested in the kinetic
148 experiments were 0, 5, 30 and 10, 20, 40 g L^{-1} , respectively. After the experiment
149 started, sub-samples were collected at 0.5, 1.0, 2.0, 4.0, 6.0, 10.0, 16.0, 24.0, 34.0,
150 48.0 h. Other experimental conditions were the same as described for the adsorption
151 experiment.

152 **4) Salt specific experiments**

153 In order to assess the specific effects of different sea salt ion compositions on
154 DIP adsorption, CaCl_2 , MgCl_2 , Na_2SO_4 , and NaHCO_3 solids were added, respectively,
155 to NaCl solution with 15 salinity to form a final concentration of 0.01 M for the added
156 salts. This led to a slight increase in the salinity by between 0.84 to 1.42. Then, the P
157 adsorption experiments were conducted with a sediment concentration of 5 g L^{-1} . In
158 order to compare influence of multiple ions on DIP adsorption, an adsorption
159 experiment with salinities at 5, 15, 30 and sediment concentration at 10 g L^{-1} was
160 conducted in seawater, NaCl solution and artificial seawater.

161 **2.3. Statistical analysis**

162 All the data of the amount of adsorbed and desorbed P was subjected to the
163 ANOVA using MIXED procedure (SAS Institute 2010) to evaluate the effects of
164 salinity, sediment concentration, and solution types with their interaction. The effects
165 were deemed significant when $P < 0.05$; differences among least square means for all
166 of the treatment pairs were identified using LSMEANS (/diff) statement (t-test). The
167 results are listed in Tables S2 and S3.

168

169 3. Results and discussion

170 3.1. Scavenging effect of suspended sediments

171 Field-based studies reported decreasing DIP concentrations with increased
172 loading of suspended sediments during flood season (Li et al., 2017; Ding et al., 2019).
173 Our laboratory experiments explain these results by showing that with an increase in
174 suspended sediment loading (S_c , g L^{-1}), the total amount of adsorbed P (Q_{ad} , μmol)
175 onto the suspended sediment generally increased (**Fig. 1a**), i.e. removing more than
176 80% of the DIP in water. The relative amount of adsorbed P normalized to the amount
177 of sediment ($q_{ad} = \mu\text{mol } Q_{ad} / \text{g sediment}$) decreased with increased suspended
178 sediment concentration (**Fig. 1a**), indicating a limitation of DIP on the adsorption
179 equilibrium.

180 The total amount of desorbed P (Q_{de} , μmol) from the suspended sediment
181 increased almost linearly with the suspended sediment concentration for
182 concentrations below 20 g L^{-1} (**Fig. 1b**). Above this level, the Q_{de} decreased slightly
183 with increased suspended sediment (**Fig. 1b**) due to re-adsorption of desorbed P to the
184 suspended sediments. This may indicate a DIP limitation to how much P can be
185 desorbed from the suspended sediments due to opposing adsorption/desorption
186 processes. Regardless, the absolute values for DIP adsorption were always higher than
187 that for P desorption. This reflects that P tends to adsorb onto suspended sediment,
188 rather than desorb from suspended sediment.

189 The normalized values of adsorbed and desorbed P (q_{ad} and q_{de}) are linearly

190 correlated with the suspended sediment concentration after logarithmic transformation
191 (**Fig. 1c**). The regression slopes were negative albeit larger than -1 (as would be
192 expected if there was no effect of sediment concentration on adsorption/desorption)
193 for all the adsorption and desorption experiments (**Table 1**). This is partly due to the
194 more effective contact between dissolved phosphorus and the suspended sediment
195 surface under the conditions of relatively low sediment loading (Cao et al., 2011).

196 The slopes of the adsorption regressions generally increased with increased
197 salinity. This indicates that the influence of suspended sediment concentration for DIP
198 adsorption is less important under low salinity (**Fig. 1c, Table 1**). On the other hand,
199 in the desorption experiments no significant difference in the slopes under different
200 salinities were observed, except for deionized water (Salinity=0) (**Fig. 1c, Table 1**).
201 This implies that for saline water the suspended sediment concentration primarily
202 controls the desorption of P. The exception for zero salinity suggests that the
203 desorption of P from the sediment suspended in freshwater is stronger than that in
204 saline water.

205 DIP absorption typically reached equilibrium within 10 hours and 80% of the
206 equilibrium within an hour (**Fig. 2**). The DIP removal percentage from aqueous phase
207 was strongly correlated ($r = 0.96 - 0.99$) with the concentration of suspended sediment,
208 reaching between 60% and 85% of DIP in water being scavenged by suspended
209 sediment when the suspended sediment concentration was 40 g L^{-1} (**Fig. 2**).
210 Increasing the salinity from 0 to 5 increased the DIP removal at equilibrium (Q_{ad}) by
211 up to 80% in the saltwater solutions (**Fig. 3**). This strong increase in removal
212 efficiency, with only slight increase in salinity, illustrates the important effect
213 increased salinity has on the sediments' scavenging of DIP in brackish estuarine
214 waters.

215 These adsorption/desorption laboratory experiments demonstrate that suspended
216 sediments in the estuarine zone have a strong capacity to efficiently and rapidly
217 remove DIP from estuarine water. This is especially important in estuaries receiving

218 discharge from gated rivers. The strong hydrodynamic force when the gate opens re-
219 suspends bed sediments, thereby episodically increases the concentration of
220 suspended sediments in the gated estuary. Increasing salinity in estuaries enhances the
221 DIP scavenging by the suspended particles. Salinity increase also causes the
222 suspended particles with the adsorbed P to flocculate into larger particles that settle
223 faster in the deeper estuarine channels (Snedden et al., 2007; Figueroa et al., 2020a,
224 2020b). The fate of the DIP absorbed to settling particles remains unknown, but it is
225 possible that the actual seaward flux of DIP, and hence total P, is overestimated by
226 disregarding the scavenging effect by suspended sediment, especially for the estuaries
227 of gated rivers.

228 3.2. Salinity effect

229 Due to the predominant inner-sphere specific chemical bonding of PO_4^{3-} to net
230 negatively charged particles, ions such as Ca^{2+} , Mg^{2+} , SO_4^{2-} and HCO_3^- actually
231 determine the overall P adsorption and desorption in estuarine water. The effect of
232 increased salinity on DIP adsorption differed between the three experimental solutions,
233 i.e. NaCl solution, seawater, and artificial seawater (**Fig. 3**). Generally, the P
234 adsorption is stronger in the artificial seawater and lower in the NaCl solution. This
235 indicates that the DIP adsorption is underestimated if NaCl solution is used and
236 overestimated if the artificial seawater is used to simulate DIP sorption in estuarine
237 waters.

238 Specifically, in the case of NaCl solution, the total adsorbed P onto the sediment
239 (Q_{ad}) at low salinity (0, 5) was larger than that at high salinity (15, 30). It was reported
240 that the structural destruction of Na^+ on particle and the competitive behavior of Cl⁻
241 for active point on sediment, could reduce the P adsorption on sediment (Zhang et al.,
242 2011; Wu et al., 2012; Nguyen et al., 2019). For artificial seawater, there is a strong
243 increase of Q_{ad} from 0 to 5 salinity, but no significant change was measured with
244 further increase in salinity from 15 to 30. For seawater, the DIP adsorption under both
245 low and high salinity conditions is lower than that under moderate salinity (i.e. 15)

246 (Fig. 3). The instable influence of salinity in seawater may be caused by coexistence
247 of cation promotion and anion restraint effects on DIP adsorption (Meng et al., 2014b).
248 Similarly, it seemed that the cation effect played a more important role with the
249 increase of salinity for the artificial seawater. This entails that the underestimation of
250 DIP adsorption for NaCl-solution-based experiments would increase with salinity, and
251 the overestimation of DIP adsorption for artificial-seawater-based experiments would
252 be larger for low and high salinity conditions.

253 The differences in DIP adsorption change with salinity among the three types of
254 saltwater solutions may conceivably be related to their ionic compositions, as
255 indicated in Fig. 4. Divalent cations, such as Mg^{2+} and Ca^{2+} , conceptually promote
256 phosphate adsorption on sediments, by forming bridge binding between negatively
257 charged surface functional groups and the phosphate anion, described above,
258 enhancing the number of DIP adsorption sites on sediments surface (Millero et al.,
259 2001; Spiteri et al., 2008). Concurrently, divalent cations promote coagulation and
260 flocculation and thereby settling of the sediments (Liu et al., 2002). Although
261 restraining effect of SO_4^{2-} on PO_4^{3-} adsorption in other studies (e.g., Flower et al.,
262 2017), SO_4^{2-} was relative to Cl⁻ observed to enhance the DIP adsorption (Fig. 4). The
263 calcite content is 2% in the sediment. It was reported that sulfate may increase calcite
264 solubility, and thereby can promote Ca^{2+} release from the sediment (Liu et al., 2012).
265 The cation Ca^{2+} can offer adsorption point for phosphate on sediment. Free Ca^{2+} will
266 precipitate with SO_4^{2-} in advance because the concentration of sulfate is higher than
267 phosphate (200:1). Calcium sulfate will then be converted into hydroxyapatite acting
268 as a continuous calcium supply pool for phosphate removal (Liu et al., 2012). This
269 pathway may explain the phenomenon, as the pH of the solution was 7.0-8.0 and the
270 Ca^{2+} content of the sediment was 28.8 g kg^{-1} (Table S1). Compared to the pure NaCl
271 solution and seawater, the HCO_3^- inhibited DIP adsorption. This may be due to the
272 competition between HCO_3^- and PO_4^{3-} for the positively charged sorption sites on
273 sediment (Miller et al., 2001; Flower et al., 2017; Nguyen et al., 2019). Compared

274 with seawater, artificial seawater has higher SO_4^{2-} (3.160 g L^{-1}) and lower HCO_3^-
275 (0.139 g L^{-1}) contents (Table SI-1), and there are no other ions except Na^+ and Cl^- in
276 the NaCl solution. This difference in sulfate could partly explain why the P adsorption
277 was highest in artificial seawater, and lowest in NaCl solution (**Fig. 3**).

278 **3.3. Combined effects on seaward P flux**

279 As described above, both salinity and suspended sediment concentration have
280 significant effect on DIP adsorption on sediment. Spatially, salinity increases and
281 suspended sediment concentration decreases with the distance from the river mouth.
282 In the estuarine environment of gated rivers, the temporal changes in salinity and
283 suspended sediment concentration are governed by the strong hydrodynamic episodes
284 related to the opening of the gates. This means that the effects of salinity and
285 suspended sediment concentration on DIP adsorption and desorption are generally
286 combined in space and time. The distribution ratio (DR) between P in suspended
287 sediment and DIP in water varied with salinity and suspended sediment concentration
288 (**Fig. 5**). In the adsorption experiment using seawater (30 salinity), the DR value
289 increased with suspended sediment concentration (**Fig. 5a**). Within the same level of
290 suspended sediment concentration, the highest DIP adsorption was achieved under
291 medium salinity condition (15). However, the influence of suspended sediment
292 concentration on DIP adsorption was larger than that of salinity (cf. **Fig. 1a**). This
293 implies that during periods with open gates, the enhanced sediment resuspension is
294 probably more important for the DIP adsorption than the antagonistic effect caused by
295 the concurrent reduced salinity of the freshwater pulse.

296 In the desorption experiment using seawater, the DR values tended to increase
297 with increasing suspended sediment concentration (**Fig. 5b**), indicating that
298 desorption decreases with increasing suspended sediment concentration. On the other
299 hand, the DR value showed a sharp decrease at low salinity (<5), implying that
300 salinity may promote P desorption under such conditions. At higher salinity (>5), the
301 DR values remained stable at low suspended sediment concentration ($< 25 \text{ g L}^{-1}$).

302 However, the DR value increased with continuous increased salinity (>5) in the
303 solutions with higher suspended sediment concentration (> 25 g L⁻¹), indicating that
304 the combined effect of salinity (> 5) and higher sediment concentration (> 25 g L⁻¹)
305 may reduce P release from sediment.

306 **3.4 Implications and perspectives**

307 Both suspended sediment concentration and salinity were found to affect P
308 adsorption and desorption processes in estuaries, though the effect of suspended
309 sediment concentration is larger than salinity. Although P desorption at high sediment
310 concentration is highest around 10 salinity, under low salinity conditions the P
311 desorption is outweighed by the role of suspended sediments. Therefore, the
312 scavenging effect for DIP plays an important role, implying that the overall DIP
313 concentration decreases through the estuarine zone.

314 DIP concentrations in the world's estuaries exhibit many different patterns
315 across the salinity gradient (**Fig. 6**). These changes from land to sea are governed by
316 many different processes from physical mixing between the riverine and oceanic end-
317 member, inputs from point sources along the salinity gradient, uptake by plants and
318 bacteria, remineralization processes and release of iron-bound phosphate. In addition
319 to these well-studied processes, we submit that absorption/desorption processes
320 should also be considered for assessing the potential impact of riverine DIP inputs on
321 the coastal biogeochemistry, at least in estuaries with hydrological control of
322 freshwater discharge.

323 The pulse-like water discharge during the open gate period promotes sediment
324 resuspension, and this will enhance adsorption of DIP onto the suspended sediment,
325 potentially reducing the DIP flux to the sea. Even though moderate (15) salinity
326 initially accelerate the P desorption, higher salinity may reduce the P desorption (**Fig.**
327 **1b**). Certainly, this effect is limited by the DIP adsorption capacity of suspended
328 sediments, and the P adsorbed on buried sediment may release back into the water
329 column. However, for high sediment concentrations the P adsorption capacity could

330 be sufficient to maintain P absorbed to particles, and the release of DIP from sediment
331 to water column is generally slow and limited within the sediment surface (López-
332 Hernández et al., 1980). Hence, burial of DIP absorbed to sediment particles could be
333 an important sink for P, thereby reducing coastal eutrophication and the potential
334 release of iron-bound phosphate.

335 Simply scaling the freshwater DIP concentration with the water flow will
336 possibly provide an overestimate of the seaward flux of DIP, especially in estuaries
337 fed by gated rivers. The discrepancy between the actual and estimated values of
338 seaward DIP flux depends on the concentration of DIP and particles in the freshwater
339 discharge. The latter is determined by the number of gates, the fluctuation in
340 hydrodynamic strength, and the gate location, as well as the physical and chemical
341 properties of estuarine sediment. In addition, particle sedimentation is probably
342 enhanced when the gates shut down the upstream water flow. This suggests that
343 alternating hydrological regimes could potentially promote sediment burial of P and
344 reduce the seaward flux of DIP.

345 Therefore, it is necessary to introduce an imperial correction model to eliminate
346 the erroneous estimates of DIP seaward flux. The imperial correction model might be
347 estuary-specific, but could be derived from the calibration of lab experimental data
348 and field monitoring data. Correspondingly, the monitoring of P in the estuary should
349 cover the salinity range from freshwater to seawater and the whole process of gate
350 open and close, also the above-mentioned influencing factors.

351

352 **4. Conclusion**

353 The seaward flux of DIP from the terrestrial environment is important for
354 controlling coastal and marine eutrophication. However, this flux is possibly
355 overestimated by not accounting for the physicochemical processes taking place in the
356 gated estuaries where DIP scavenging by suspended sediment is enhanced. Increasing
357 salinity generally enhances the adsorption, and in addition, polyhaline conditions may

358 reduce P desorption. For estuaries fed by gated rivers, the seaward P flux may be
359 limited during period with open sluices due to enhanced P scavenging by higher
360 sediment resuspension, even though the salinity is lower.

361

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364

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Table 1. The linear fitting parameters of $\ln q_{ad}$ and $\ln q_{de}$ with $\ln Sc$.

Salinity		Deionized water		Seawater		NaCl solution		
		0	5	15	30	5	15	30
Adsorption	slope	-0.37	-0.54	-0.53	-0.23	-0.50	-0.44	-0.12
	intercept	1.33	2.04	2.09	1.05	1.66	1.05	0.24
	R ²	0.97	0.99	0.99	0.78	0.99	0.76	0.65
	p	0.002	0.0002	0.0003	0.05	0.0006	0.05	0.1
Desorption	slope	-0.89	-0.67	-0.66	-0.59	-0.69	-0.70	-0.61
	intercept	0.92	-0.23	0.29	-0.51	0.59	0.49	0.37
	R ²	0.99	0.97	0.99	0.99	0.98	0.97	0.96
	p	0.007	0.02	0.001	0.0003	0.01	0.01	0.02

Figure Captions

Fig. 1. The total amount of adsorbed and desorbed phosphorus (Q , $\mu\text{mol g}^{-1}$) (a), the amount of adsorbed and desorbed phosphorus for unit sediment (q , $\mu\text{mol g}^{-1}$) (b), and the linear relationships between the natural log of adsorbed/desorbed P and natural log of suspended sediment concentration (S_c , g L^{-1}) (c) under different suspended sediment concentrations and salinities, using both seawater and NaCl solutions. The error bars mark one standard deviation of the duplicates. The solid and dashed lines denote adsorption and desorption, respectively

Fig. 2. DIP removal percentage over time from the kinetic experiment for different suspended sediment concentrations (S_c) at different salinities of seawater solution.

Fig. 3. Comparison of P adsorption capacity of 10 g L^{-1} suspended sediment for different salinities of NaCl solution, artificial seawater, and seawater. Q_{ad} (a), the total amount of adsorbed P.

Fig. 4. Influence of Ca^{2+} , Mg^{2+} , HCO_3^- and SO_4^{2-} on phosphorus adsorption in NaCl solution (salinity = 15, suspended sediment concentration = 5 g L^{-1}). The concentration of the added ions is 0.01 M. Pure seawater and NaCl solution without ion addition are set as the control.

Fig. 5. The combined effects of salinity and sediment concentration in seawater on P distribution ratio (DR) between P adsorbed to sediment surface and DIP in the water column. Note that the distribution ratio is not the same as equilibrium distribution coefficient (K_d). The data were interpolated by the Kriging method.

Fig. 6. DIP concentrations along the salinity gradient for different estuaries worldwide (De Jonge et al., 1989; Upchurch et al., 1974; Ballagh et al., 2019; Meng et al., 2015; Xu et al., 2015; Lin et al., 2012; Pamplona et al., 2013).

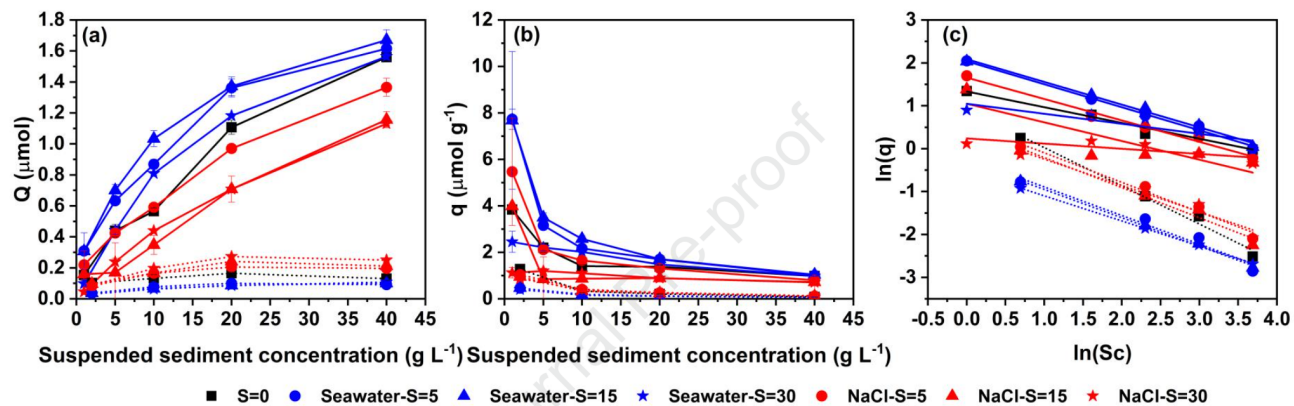


Fig. 1.

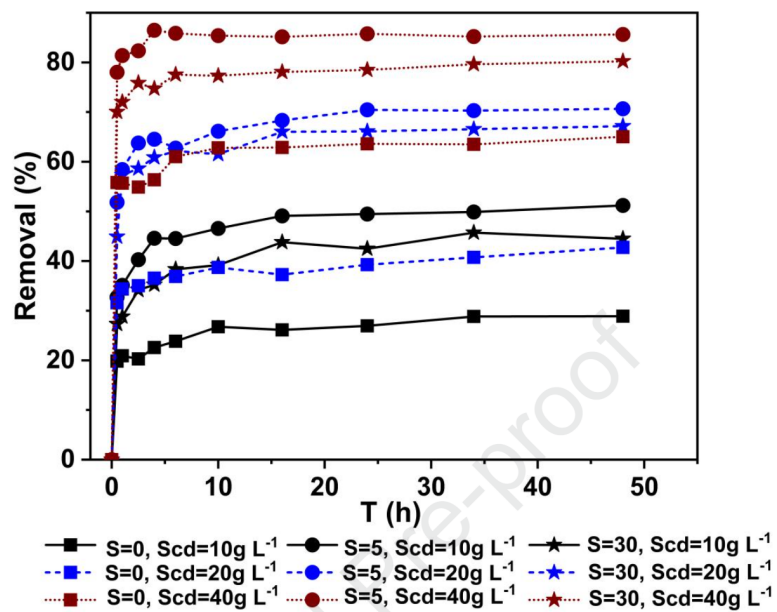


Fig. 2.

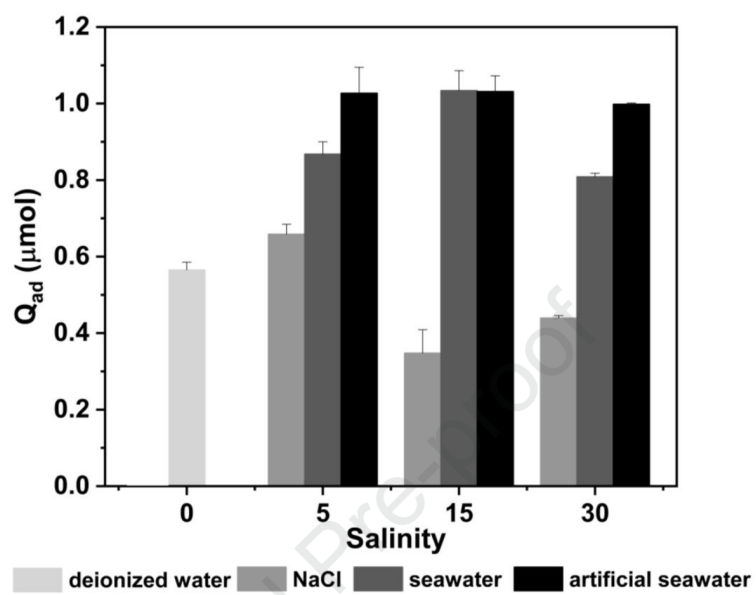


Fig. 3.

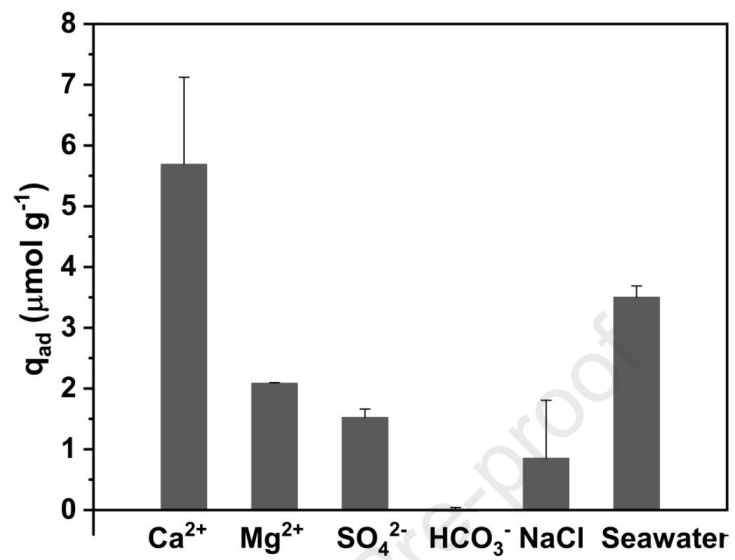


Fig. 4.

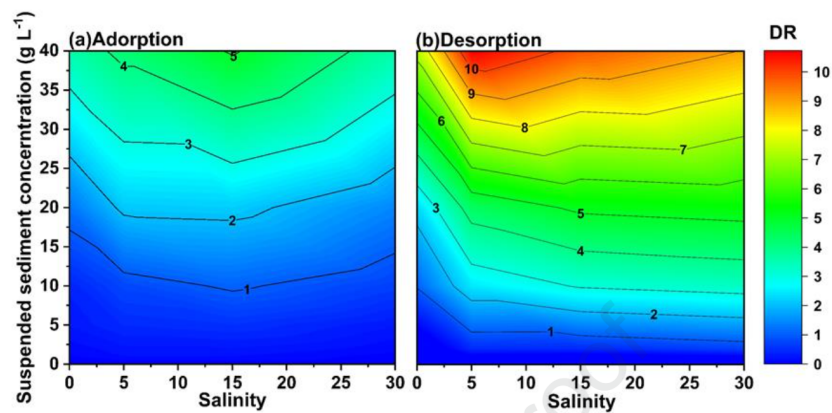


Fig. 5.

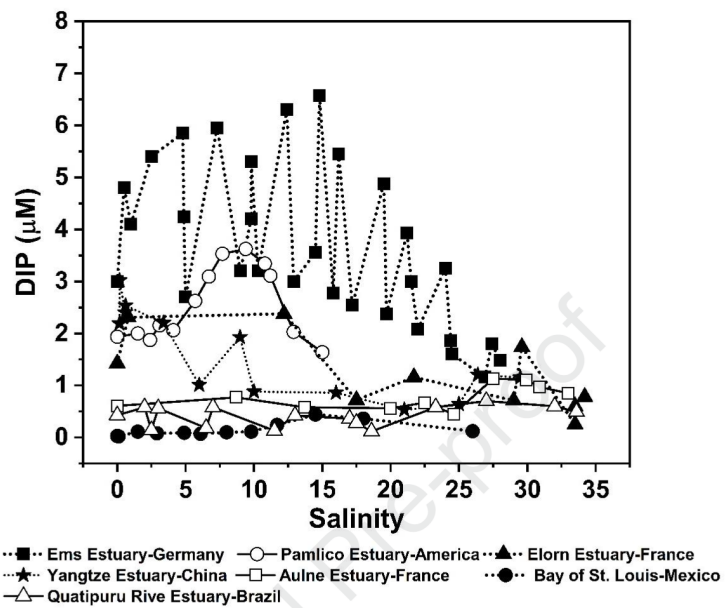


Fig. 6.

Highlights

1. DIP scavenging by suspended sediment is prevalent for gated estuaries.
2. The increasing of salinity can enhance the DIP adsorption on sediments.
3. Riverine flux of dissolved phosphorus to the coastal sea may be overestimated.
4. Not-real-seawater-based experiments may mislead the DIP flux estimation.

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Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

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