Predicting the proximity to macroscopic failure using local strain populations from dynamic in situ X-ray tomography triaxial compression experiments on rocks

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Highlights

- Machine learning successfully predicts the proximity to failure from strain fields.
- Dilatational strain predicts the proximity to failure better than shear strain.
- The 25th-75th percentile values of the strain population best predict failure proximity.

1 Abstract

2 Predicting the proximity of large-scale dynamic failure is a critical concern in the engineering 3 and geophysical sciences. Here we use evolving contractive, dilatational, and shear strain 4 deformation preceding failure in dynamic X-ray tomography experiments to examine which 5 strain components best predict the proximity to failure. We develop machine learning models 6 to predict the proximity to failure using time series of three-dimensional local incremental 7 strain tensor fields acquired in rock deformation experiments under stress conditions of the 8 upper crust. Three-dimensional scans acquired in situ throughout triaxial compression 9 experiments provide a distribution of density contrasts from which we estimate the three-10 dimensional incremental strain that accumulates between each scan acquisition. Training 11 machine learning models on multiple experiments of six rock types provides suites of feature 12 importance that indicate the predictive power of each feature. Comparing the average 13 importance of groups of features that include information about each strain component 14 quantifies the ability of the contractive, dilatational and shear strain to predict the proximity 15 of macroscopic failure. A total of 24 models of four machine learning algorithms with six 16 rock types indicate that 1) the dilatational strain provides the best predictive power of the strain components, and 2) the intermediate values (25th-75th percentile) of the strain 17 18 population provide the best predictive power of the statistics of the strain populations. In 19 addition, the success of the predictions of models trained on one rock type and tested on other 20 rock types quantifies the similarities and differences of the precursory strain accumulation 21 process in the six rock types. These similarities suggest the potential existence of a unified 22 theory of brittle rock deformation for a range of rock types.

23

24 Keywords

25 strain localization; tomography; machine learning; triaxial compression

26 1. Introduction

27 Predicting the timing of dynamic catastrophic failures is a fundamental problem for 28 society and earthquake science. This question could be more straightforward to address if 29 large faults were planar and surrounded by a purely elastic and homogeneous solid where 30 strain energy only accumulates in the time between the episodic failures of the large faults. 31 However, seismic and geodetic observations document that crustal volumes experience 32 ongoing seismic and aseismic failure on a heterogeneous population of faults with various 33 sizes, geometries and failure mechanisms (e.g., Scholz, 1990; Ben-Zion, 2008; Hauksson et 34 al., 2012).

Under triaxial compression conditions, permanent deformation preceding macroscopic
failure manifests by the decreasing slope of the axial strain-stress curve following a nominally
linear phase, along with a corresponding reduction of the effective moduli toward
macroscopic failure, and increasing transverse or radial deformation (e.g., *Brace et al.*, 1966; *Scholz et al.*, 1973; *Lyakhovsky et al.*, 1997; *Hamiel et al.*, 2004). Distributed dilation within
the rock volume preceding dynamic rupture helps explain the emergence of these and other
precursory signals.

42 The dilatancy-diffusion model posits that as fractures propagate and open in the presence 43 of fluids in the crust, the increased pore space allows fluid to diffuse, producing larger ratios 44 of fluid to solid volume (e.g., Nur, 1974). The increased ratio of fluid (air or liquid) to solid 45 may reduce the P-wave velocity and decrease the effective elastic modulus, and dilatancy-46 enhanced fluid flow may change local hydrologic levels. Some field observations are 47 consistent with these expectations (e.g., Frank, 1965; Whitcomb et al., 1973; Aggarwal et al., 48 1973; Roeloffs, 1988), but the general validity of these observations remain controversial 49 (e.g., McEvilly & Johnson, 1974; Bolt, 1977; Haase et al., 1995). Observations of 50 macroscopic volumetric strain during triaxial compression experiments indicate that rocks

51 tend to dilate before macroscopic failure under confining stresses representative of the upper 52 crust (e.g., *Brace et al.*, 1966). Quantifying the behavior of rocks during the approach to 53 macroscopic failure is a critical step toward constraining the precursory signals that reveal the 54 proximity of the next large earthquake.

55 A key assumption of the earthquake preparation process is that both the dilatational and 56 shear deformation evolve toward failure, and that both modes of deformation contain 57 information critical for predicting the timing of rupture. For example, the second invariant of 58 the strain rate tensor, which includes the volumetric and deviatoric strain tensor components, 59 is often used as a proxy for long term seismic hazard (e.g., Kreemer et al., 2003). Earthquakes 60 can rupture in a combination of modes, and so the reported accelerated seismic release rates before mainshock events (e.g., Mogi, 1969; Papazachos, 1973; Sykes & Jaumé, 1990) may 61 62 involve different modes of deformation. However, increased rates of foreshocks surrounding 63 normal faults, relative to thrust and strike-slip faults (Abercrombie & Mori, 1996), suggests 64 that extensional deformation may provide less ambiguous signals than shear deformation 65 before rupture. Similarly, the observed acceleration of the total fracture volume in X-ray 66 tomography experiments suggests a process consistent with a critical phase transition (Renard 67 et al., 2018; Kandula et al., 2019), but this tracking of the fracture volume does not provide 68 information about the evolution of shear deformation.

Time series of 3D incremental strain component fields throughout X-ray tomography experiments (e.g., **Figure** 1) suggest that both the dilatational and shear strain components accelerate toward failure in some experiments of some rock types (e.g., *Renard et al.*, 2019). The dilatational strain, as measured with the mean of the local incremental strain population, tends to accelerate in more experiments than the shear strain (*Renard et al.*, 2019; *McBeck et al.*, 2018, 2019, 2020). The consistency of the dilatational strain evolution suggests that shear strain may not provide information crucial for predicting the timing of macroscopic failure. In

this case, effective predictive analyses could concentrate on detecting the precursory signalsindicative of dilatancy.

78 To determine which strain components are most likely to improve the ability of predicting 79 the proximity to system-size failure, we analyze the time series of 3D local incremental strain 80 fields captured in X-ray tomography triaxial compression experiments (e.g., Figure 1). Here, 81 we characterize the proximity to failure by the ratio between the differential stress at a given 82 time, and the differential stress at the macroscopic failure of the sample. The incremental 83 strain fields are calculated throughout twelve in situ dynamic X-ray tomography triaxial 84 compression deformation experiments on six rock types (Figure 2, Table S1). Machine 85 learning models then use the statistics of the local contraction, dilation and shear strain 86 populations to predict the proximity to failure (Table S2). We examine the feature importance 87 and Shapley Additive Explanation (SHAP) values to determine which strain components and 88 statistics have the best predictive power of the proximity to macroscopic failure. To assess 89 differences in the predictive power of these strain components and statistics among different 90 rock types, we train six unique models on datasets from two experiments on the same rock 91 type (Table S1). To quantify the similarity of the precursory strain accumulation processes in 92 these rock types, we test the success of the models with data from the same rock type used to 93 train the model, and with data from different rock types. High degrees of success between 94 models trained and tested on different rock types may reflect the similarity of the strain 95 accumulation process in these rock types. To increase the robustness of the results, we repeat 96 the analysis using four machine learning approaches, including random forest and gradient 97 boosting (XGBoost) methods. This suite of analyses indicates that for the six analyzed rock 98 types, the intermediate values of the dilatational strain populations provide the best predictive 99 power of the proximity to macroscopic failure.

101 **2. Methods**

102 **2.1. Experimental design and data**

103 We deform rocks in an X-ray transparent deformation apparatus installed on beamline 104 ID19 at the European Synchrotron and Radiation Facility. In each experiment, we increase the 105 axial stress in steps of 0.5-5 MPa under constant confining stress between 5-35 MPa until the 106 sample fails macroscopically (Table S3). At each stress step, we acquire a 3D X-ray 107 tomogram at 6.5 µm per voxel-side resolution while the sample is under constant stress 108 conditions inside the apparatus. Each scan requires about 2 minutes, and the final scan 109 immediately precedes macroscopic failure of the rock. McBeck et al. (2020) describe in detail 110 the experimental conditions and method of calculating the strain tensor fields used in the 111 current analysis, so we only briefly describe these topics here. We calculate the differential stress, $\sigma_D = \sigma_1 - P_c$, where σ_1 is the axial stress and P_c is the confining pressure. The 112 113 differential stress at failure is noted σ_F . 114 From the 3D tomograms of the 12 experiments, we perform digital volume correlation 115 (DVC) in ten intervals of each experiment (e.g., Figure 2), following the approach of *McBeck* 116 et al. (2020) with the code TomoWarp2 (Tudisco et al., 2017). Each interval is separated by 117 approximately equal increments of cumulative macroscopic axial strain. The DVC 118 calculations provide the 3D displacement fields between each scan acquisition, and thus the 119 incremental strain tensor components. The positive and negative divergence represents the dilatational and contractive volumetric strains, respectively, and the magnitude of the angular 120 121 velocity of the curl represents the shear strain.

122 **2.2. Feature extraction**

We train the models using data from all the experiments with the same rock type, producing six unique models of each rock type (**Table** S1). To extract information about the 3D incremental strain field, we subdivide each strain field into a grid of subvolumes that are

126 fixed in space from which we report statistics of the strain population. The 27 features of the 127 data are a combination of one of the three strain components (dilation, contraction, shear 128 strain), and a measurement of the population of these strain components within the subvolume 129 (Table S2). Each subvolume in the cubic grid has a side length of 0.5 mm, about four times 130 the DVC spatial resolution (0.13 mm), and is fixed in space. Therefore, each subvolume includes about 4³ measurements of each of the three strain components. The strain 131 132 components include contractive, dilatational and shear strains, as measured with the negative 133 divergence, positive divergence, and magnitude of curl of the displacement field. The statistics include the 90th, 75th, 50th, 25th, and 10th percentile, mean, standard deviation and 134 135 sum of the strain population, and the number of measurements within a subvolume. For example, one feature is the 90th percentile of the contraction values within a subvolume of the 136 137 strain field. The number of measurements within a subvolume is the total number of strain 138 values within the subvolume, which tends to increase for dilatational strains and decrease for 139 contractive strains with loading. This number will be the same for the shear strain values 140 throughout loading, and equal to about 4³. We extract the number of measurements within a 141 subvolume because previous work shows that the volume of rock that the dilatational and contractive strains occupy can accelerate toward failure (e.g., Renard et al., 2019), and so 142 143 may signal the approach of macroscopic failure. Thus one subvolume produces one sample 144 for the machine learning models, which contains 27 features. Testing the influence of using a 145 smaller sub-volume size (0.2 mm) revealed only insignificant differences in the results.

146 **2.3**

2.3. Machine learning methods

We employ two machine learning methods: random forest classification (e.g., *Breiman*,
2001) and XGBoost (i.e., extreme gradient boosting) classification (e.g., *Chen & Guestrin*,
2016). We designed this analysis as a classification problem similar to time-dependent

150 seismic hazard models that predict the rate of event occurrence (e.g., *Tiampo & Shcherbakov*,

151 2012). We synthesize results from the random forest and XGBoost methods to increase the 152 robustness of the conclusions. Both methods utilize an ensemble of decision trees as the base 153 estimator. **Text** S1 describes differences in the random forest and XGBoost methods. For 154 each machine learning method, we perform a grid search over the hyperparameter space to 155 find the appropriate set of hyperparameters for each model (*Géron*, 2017).

156 The models predict the proximity to system size failure using the macroscopic differential 157 stress applied to the rock at the time that the strain field was calculated, and the differential 158 stress at failure, σ_F . We use the differential stress an indicator of the distance until failure 159 because previous experimental studies have demonstrated that fractures develop in rocks 160 preceding macroscopic compressive failure as a critical phase transition relative to the 161 macroscopic σ_F (e.g., *Renard et al.*, 2018; *Kandula et al.*, 2019).

162 We use two different classification schemes for each of the two machine learning 163 methods, producing four machine learning approaches. We predict either two classes or four 164 classes of the proximity to failure. For the four prediction classes, we predict whether the rock 165 is experiencing <25%, 25-50%, 50-75% or >75% of σ_F . For the two prediction classes, we 166 predict whether the rock is experiencing <50%, or >50% of σ_F . For example, if the rock is under 10 MPa differential stress, σ_D , and fails at σ_F =100 MPa, then σ_D/σ_F = 0.10 or 10%, so 167 168 this distance to failure is in the first prediction class in both the four- (<25%) and two-class 169 (<50%) prediction models. This classification approach enables comparing the model 170 accuracy to the accuracy expected from the random probability of selecting the correct 171 answer, providing a systematic means of quantifying the prediction success. The accuracy is 172 the number of correct answers out of the total number of correct and incorrect answers (e.g., 173 Müller & Guido, 2016).

174 The time series of incremental strain fields are calculated with approximately equal 175 increments of macroscopic axial strain, so predicting the proximity to failure using the macroscopic differential stress may result in different numbers of outcomes (stress classes)
depending on the non-linearity of the stress-strain relationship. Using more than four classes
results in several experiments having a different number of outcomes (predicted stress
classes). Such differences in the expected random probability between experiments reduces
the likelihood of achieving high transfer accuracies, when models are trained on data from
one rock type and tested with data from another rock type. For this reason, we do not use
more than four prediction classes.

183 With the 50% partitioning of the experiments into two stress classes, all of the 184 experiments have two classes of predictions. Consequently, the random probability of 185 selecting the correct answer is 1/2. With the 25% partitioning of the experiments into four 186 stress classes, all but two of the experiments have four stress classes. The shale experiments 187 include longer yielding phases preceding macroscopic failure than the other rock types, so 188 they only have three stress classes. Consequently, whereas the shale experiments produce a 189 random probability of 1/3, all the other experiments have four prediction classes with a 190 random probability of 1/4. We counted the number of outcomes in each class in the testing 191 datasets to ensure that the expected probabilities are the same as those in the testing datasets. 192 After counting the number of outcomes in each class, we found that the mean accuracies 193 across all the classes expected from a random guess in the testing datasets were the same as 194 those in the training datasets.

To determine which strain components and statistics signal the approach to macroscopic failure, we examine the suites of feature importance of the trained models. The feature importance quantifies the average information gain produced by encountering a feature. The magnitude of improvement in the model prediction associated with a node and its corresponding information is defined as the node impurity. The decrease in node impurity weighted by the probability of reaching the node produces the feature importance (e.g.,

Müller & Guido, 2016). Higher feature importance indicates that the feature has more
influence in fitting the data, so we interpret the importance as indicating the relative
predictive power of features.

204 In addition to the feature importance, we use the Shapley Additive Explanation (SHAP) 205 values to examine the impact of individual features on the feature space (Lundberg & Lee, 206 2017). SHAP values indicate the contribution of each feature on the prediction of the average 207 (i.e., a completely featureless naïve model). They may explain single predictions for single 208 measurements, and also for the collection of predictions in a data set (i.e., stress classes). 209 SHAP values for a specific feature S are calculated by comparing results of models that do 210 and do not include feature S. The approximate effect on all predictions may then be estimated 211 by finding the mean SHAP value over the samples in the training data.

212

213 **3. Results**

214 **3.1. Transfer and non-transfer success**

215 First, we examine the success of the models when trained and tested on the same rock 216 type, and with distinct datasets. In particular, we trained six unique models using 80% of the 217 data from experiments on the same rock type: sandstone, basalt, monzonite, granite, shale, 218 and limestone (Figure 3). We then tested the accuracy of the models with the remaining 20% 219 of the data from each group of experiments (i.e., non-transfer accuracy), and with data from 220 the other groups of experiments (i.e., transfer accuracy) (Figures 4, S2, S3, S4). We divided 221 the data into training and testing datasets by random selection. Due to potential auto-222 correlation between the data in the training and testing datasets, the non-transfer accuracies 223 may be higher than the transfer accuracies.

The four machine learning approaches include the two schemes of prediction with two and four classes of prediction, and two algorithms of the random forest and XGBoost

226 implementations. The four machine learning approaches and six rock types produce 24 227 models. The 12 models that predict two classes have a higher range of accuracy (0.76-0.89) 228 than the 12 models that predict four classes (0.58-0.83), as expected from the random 229 probability of selecting the correct answers of 0.50 and 0.25 for the two and four class 230 prediction schemes, respectively (Figure 3). The XGBoost models tend to perform better, in 231 terms of accuracy, than the random forest models. To determine which part of the strain 232 accumulation process these models most successfully describe, we extract the non-transfer 233 precision and recall of each prediction class for the 24 rock type models (Figure S2). In 234 general, the precision scores indicate that the model predictions are similarly successful 235 across all the prediction classes. The recall scores indicate that the models are more correct 236 near the onset of loading.

237 We next assess the ability of the models to predict the proximity to failure with data from 238 experiments on rock types different than those used to train the models (i.e., transfer 239 accuracy). The non-transfer accuracy reflects the success of the model in predicting the 240 outcome with data from the same rock type. The transfer accuracy reflects the success of the 241 model in predicting the outcome with data of different rock types. High transfer accuracies 242 may indicate that the rock type used in the testing dataset fails with similar a strain 243 accumulation process as the rock type used in the training data. Note that higher non-transfer 244 accuracies may arise if samples used in the training dataset are moderately-highly correlated 245 to samples used in the testing dataset. We show below that this caveat has minimal impact on 246 the model success.

To compare the similarity of the strain accumulation processes in these rock types, we synthesize the accuracies of the models trained and tested with data from the 36 rock type pairs in each of the four machine learning approaches. First, we find the accuracy of the 36 rock type pairs for the four machine learning approaches (**Figures** S3, S4). Then we calculate

the difference between the accuracy and the random probability of selecting the correct answer, which is 0.50 and 0.25 for the two and four prediction classes, respectively. Then we calculate the mean of this score from the four approaches for the 36 combinations of training and testing rock types (**Figure** 4).

255 We expect that certain pairs of rock types will deform with similar strain accumulation 256 processes. For example, we expect that sandstone and basalt, monzonite and granite, and 257 shale and limestone will host similar strain accumulation processes because previous work 258 has observed similar deformation mechanisms operating in these rocks (e.g., Reches & 259 Lockner, 1994; Menéndez et al., 1996; Lockner, 1998; Katz & Reches, 2004; Zhu et al., 260 2010). The results are consistent with these expectations for the sandstone and basalt 261 experiments, and shale and limestone experiments, but not the monzonite and granite 262 experiments (Figure 4). In particular, the 1) sandstone and basalt, and 2) shale and limestone 263 datasets have high transfer accuracy, as measured with the mean accuracy above the random 264 probability, but the 3) monzonite and granite transfer accuracies are relatively lower. Whereas 265 the difference in the non-transfer and transfer accuracies is 0.10-0.16 for the sandstone and 266 basalt, and <0.15 for the limestone and shale, this difference is 0.15-0.20 for the monzonite 267 and granite. Monzonite and granite are low porosity crystalline rocks comprised of quartz and 268 feldspar, so we would expect them to fail in similar strain accumulation processes. In contrast 269 to expectations, the transfer accuracy is >0.15 lower than the non-transfer accuracy for these rocks. The general similarity between the range of non-transfer and transfer accuracies 270 271 suggest that the influence of auto-correlation on the non-transfer model performance was 272 minimal.

273 **3.2.** Predictive power of strain components

The importance of the features decays from the most to the least important in all 24 models of the six rock types and four machine learning approaches (**Figure S5**). The feature

276 importance values are measured relative to other features in the model (e.g., Müller & Guido, 277 2016). However, the similar magnitudes of the feature importance across the models indicate 278 that we may define thresholds in order to identify the highly important features of each model 279 to then compare to other models. Consequently, we examine the average importance of 280 features that have >25%, >50% and >75% importance of the maximum feature importance 281 (Figure 5). To ensure that the magnitudes of the suites of feature importance of different 282 models do not influence our results, we also examine the feature importance normalized by 283 the maximum feature importance of each model. This normalization yields the same 284 conclusions (described below) as analyzing the importance without normalization.

285 To assess which strain components (contraction, dilation, and shear strain) provide the 286 best predictive power of the proximity to macroscopic failure, we examine the importance of 287 features, grouped by strain component, that are identified as highly important. With the least 288 restrictive inclusion of features (>25% of the maximum importance) (Figure 5), features that 289 include information about all of the strain components are identified as important for each 290 rock type. With the intermediate inclusion of features (>50%), a lower number of rock type 291 models depend on the contraction and shear strain than with the >25% threshold. With the 292 most restrictive inclusion of features (>75%), features that include information about the 293 dilation are considered important for five of six of the rock types. Only the limestone models 294 rely on information about the contraction, dilation and shear strain to predict failure with this 295 75% thresholding.

To assess which statistics of the strain populations provide the best predictive power about the proximity to macroscopic failure, we now examine the average importance of features that use these statistics. For each threshold of importance (25-75%), features using the sum, mean, and 25th-75th percentile of the strain population within each sampling subvolume have the highest importance (**Figure 6**, **Figure S6**). Features that use the extreme values of the strain

population, including the highest (90th percentile) and lowest (10th percentile) extremes, do
not rank as high in importance as the intermediate values. Previous machine learning analyses
have used the variance, or standard deviation, of acoustic emissions to detect approaching
failure in friction experiments (e.g., *Rouet-Leduc et al.*, 2017; *Hulbert et al.*, 2019). In our
analysis, the standard deviation of the strain components in a subvolume does not provide
significant predictive power of the timing of macroscopic failure (Figure 6, Figure S6).
To test the robustness of the results gleaned from the feature importance, we assess the

308 predictive power of the features using their Shapley Additive Explanation (SHAP) values 309 (*Lundberg & Lee*, 2017). Consistent with the analysis of the suites of feature importance, this 310 analysis suggests that features that include information about the dilation, and intermediate 311 and global values of the strain population provide the best predictive power (**Figure** S7).

To further explore the robustness of the results derived from the feature importance and SHAP values, we train and test new models that only use features that include information about the dilation, rather than all three strain components (**Figure** S8). The accuracies of the models decrease by 0.01-0.10 when they only rely on the dilation. Both the feature importance and SHAP values indicate that the statistics that provide the highest predictive power tend to measure the intermediate values of the dilation population, including the mean, median, sum, and number of values within a subvolume.

To help understand why the models select these sets of features as important, we examine the evolution of the magnitude of the strain components throughout three characteristic experiments (**Figure** 7). With the highest threshold of importance (>75%), these experiments depend on features with information about the dilation (sandstone and monzonite), and contraction (limestone), respectively (**Figure** 7). These strain components tend to have higher mean values of the total population at a stress step than the strain components that are not identified as important. In the sandstone experiment (**Figure** 7a), the mean of the population

of dilation values in each DVC calculation exceeds the mean of the shear strain and contraction after 0.6 normalized macroscopic axial strain. This trend holds for the dilation and contraction in the monzonite and limestone experiments, respectively, for at least 90% of the macroscopic axial strain. Interestingly, the trend of the dilation relative to axial strain shows a clear acceleration toward failure in the sandstone experiment (Figure 7a). However, the mean values of the strain components identified as important in the other experiments do not increase toward failure (Figure 7b-c).

333 We would expect that the strain components that help predict the proximity to 334 macroscopic failure evolve over increasing differential stress (and axial strain) such that the 335 strain population at low differential stress would be different from the strain population at 336 higher differential stress. However, this expected trend is only evident in the sandstone 337 experiment, at least by tracking the mean of the strain population throughout the rock core. 338 Although the mean values of the dilation and contraction do not systematically increase in the 339 monzonite and limestone experiments, the machine learning analyses reveal that these strain 340 components can help predict the proximity to macroscopic failure. Thus, the machine learning 341 analyses provide insights into the predictive power of these strain components that are not 342 evident from this visual inspection of the evolution of the strain components.

343

344 **4.** Discussion

345 4.1. Quantifying the similarity of the precursory strain process in rocks

346 Previous studies characterized the micromechanics of deformation preceding macroscopic
347 failure for different rock types (e.g., *Reches & Lockner*, 1994; *Menéndez et al.*, 1996;

- 348 Lockner, 1998; Katz & Reches, 2004; Zhu et al., 2010). The deformation mechanisms
- 349 developed for different rock types, such as pore collapse and Hertzian fracturing in porous
- 350 sedimentary rocks (e.g., *Menéndez et al.*, 1996; *Zhu et al.*, 2010), have not been applied to

other (e.g., low porosity) rock because the mechanisms generally depend on the
microstructure of the rock. Machine learning analyses such as those employed here can clarify
which rock types, from sandstone to granite, have similar strain accumulation processes
preceding macroscopic failure. The similarity of processes may highlight rock types that share
a dominant deformation mechanism preceding dynamic rupture.

356 We may expect that 1) sandstone and basalt, 2) monzonite and granite, and 3) shale and 357 limestone have similar dominant strain accumulation processes. Fontainebleau sandstone and 358 Mt. Etna basalt are both porous rocks associated with failure mechanisms that include stress 359 concentrations at the edges of pores and/or grains that promote fracture nucleation, and the 360 potential of fracture termination at the edges of pores and/or grains (e.g., Stanchits et al., 361 2006; Zhu et al., 2010). Monzonite and granite are both low porosity crystalline rocks 362 comprised of similar minerals with lower degrees of local mechanical heterogeneity (i.e., 363 compressive strength and stiffness) than sedimentary rocks with higher porosity. Shale and 364 limestone are associated with compactive failure mechanisms, such as pore-collapse and 365 compaction bands, and accommodating larger magnitudes of inelastic deformation prior to 366 macroscopic failure (e.g., Railsback, 1993; Baud et al., 2000; McBeck et al., 2018; Huang et 367 al., 2019). However, both shale and limestone can exhibit macroscopic dilation under triaxial 368 compression deformation (e.g., Baud et al., 2000).

The high transfer accuracies of the developed machine learning models suggest fundamental similarities of the strain accumulation process in the examined rock types (Figure 4). Some trends in the transfer accuracies are expected: the 1) sandstone and basalt, and 2) shale and limestone models and datasets produce high transfer accuracies (Figure 4). Some trends in the transfer accuracies are unexpected. For example, the granite and monzonite models and datasets do not have high transfer accuracy. One explanation of this difference could be the varying ranges of confining stresses applied to the granite (5-10 MPa)

376 and monzonite (25-35 MPa) experiments. Another difference between Westerly granite and 377 monzonite is their mean grain sizes. The Westerly granite used here has a grain size in the 378 range 100-200 µm, whereas monzonite has a larger mean grain size of 450 µm (Aben et al., 379 2016). This difference in grain sizes may cause fractures to be impeded at grain boundaries in 380 the granite at higher rates than in the monzonite. This higher rate of impedance could result in 381 more distributed fracture networks in the granite than the monzonite, which may produce 382 differing degrees of strain localization preceding dynamic failure in the rocks. These varying 383 degrees of strain localization, in turn, may produce the lower transfer accuracies of models 384 trained and tested with the local strain data of the monzonite and granite experiments.

The high transfer accuracies, and related similarity of the suites of feature importance, for the 1) sandstone and basalt, 2) shale and limestone, and 3) monzonite, shale and limestone datasets suggest that a unified general theory may be able to describe the strain accumulation process in these groups of rock types.

4.2. Predicting the proximity to failure with the intermediate dilatational strain

390 The ranking of the predictive importance among the strain components and statistics is 391 similar for the four machine learning approaches and six rock types (Figures 6, S7, S8). This 392 consistency provides confidence in the overall conclusion that the intermediate values of the 393 dilatational strain provide the best predictive power of the strain components and statistics. 394 We now tentatively extend the observations made from the experimental data to implications 395 for detecting the earthquake preparation process. Although earthquakes have been idealized as 396 dominantly frictional phenomena, interseismic strengthening of fault zones and surrounding 397 crustal volume may increase both the cohesive and frictional strength (e.g., Qiu et al., 2020). 398 This strengthening may then cause earthquakes to involve the fracture of partially healed fault 399 zone material and the surrounding crust, and then subsequent relocalization followed by 400 frictional sliding. Consequently, macroscopic failure under triaxial compression performed in

401 these experiments may represent the relocalization deformation stage in a partially healed
402 cohesive fault zone prior to an earthquake, before frictional slip along a well-defined principal
403 slip zone begins.

404 The conclusion that the intermediate values of the local strain have greater predictive 405 power than the extreme values of the local strain suggests that in order to predict seismic 406 hazard, it is important to examine data recorded in the volume of crust adjacent to the main 407 faults that localize the highest magnitudes of strain. Such analyses should consider the 408 seismicity and deformation in crustal volumes around major faults to estimate the strain value within the 25th-75th percentile of the population, rather than the extreme values that eventually 409 410 localize along the principal slip zones. The conclusion that the dilatational strain provides the 411 highest predictive power, relative to the contraction and shear strain, for a diverse range of 412 rock types and four machine learning approaches, suggests that previous methods that track 413 the influence of dilatancy (e.g., Whitcomb et al., 1973) may be able to constrain the timing of 414 earthquake rupture (e.g., Figure 8).

415

416 **5.** Conclusions

417 Proposed theories of the micromechanical deformation of rock leading to macroscopic 418 failure tend to depend on the examined rock type (e.g., Katz & Reches, 2004). Developing a 419 general method that predicts the timing of dynamic rupture for a wide variety of rock types, 420 from porous sandstone to lower porosity granite, would represent a fundamental advance in 421 rock physics. Moreover, an improved ability to forecast system-size ruptures would have 422 significant societal benefits. Our machine learning analyses suggest that groups of rock types 423 have similar strain accumulation processes preceding macroscopic failure. The similarity of 424 these processes may highlight a dominant mechanism of deformation preceding system-size 425 rupture. The models suggest that the precursory strain accumulation process is similar in 1)

426 sandstone and basalt, 2) shale and limestone, and 3) monzonite, shale and limestone (Figures
427 4, S4). These results indicate that a wide range of rock types accumulate precursory strain
428 with similar expressions before macroscopic failure.

429 Our suite of machine learning analyses suggests that the intermediate values (25th-75th 430 percentile) of the local dilatational strain provide the highest predictive power of the 431 proximity to macroscopic failure for a diverse group of rock types including basalt, granite 432 and shale under the stress conditions of the upper crust (e.g., Figure 6). This result suggests 433 that tracking deformational processes that produce dilatancy (i.e., opening fractures), and the 434 corresponding geophysical signatures such as variations in seismic velocities and density, can 435 provide accurate estimates of the onset of macroscopic failure without relying on explicit 436 information about the shear deformation. The results of this work provide incentives to 437 enhance monitoring efforts focused on effects of dilatancy in crustal volumes around large 438 faults.

439

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449 **References**

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573 **Figure 1**

Characteristic snapshots of the strain fields in 574 575 a) sandstone, b) basalt, c) monzonite and d) 576 limestone under lower and higher differential 577 stress, σ_D . The dots show where the value of the 578 contraction (black), dilation (blue) and shear 579 strain (red) is greater than the 90th percentile of 580 the population of the given strain component. 581 The localization of strain components evolves from lower to higher σ_D , suggesting that the 582 583 strain fields may indicate the proximity to 584 macroscopic failure.





586 Figure 2

Loading history and conditions of digital volume correlation calculations for experiments on sandstone (a), basalt (b), monzonite (c), granite (d), shale (e), and limestone (f). Black dots show the differential stress and axial strain when each tomogram (i.e., 3D density field) was acquired. Red lines show the loading conditions of the tomograms used in each digital volume correlation calculation. In each experiment, macroscopic failure occurs <0.1 MPa after the maximum differential stress.



593

594 Figure 3

595 Accuracy of the six rock type models for the four machine learning approaches, including 596 predicting two and four classes of the proximity to failure, with the random forest and XGBoost 597 algorithms. The predictions of two classes have a higher range of accuracy (0.76-0.89) than the 598 predictions of four classes (0.58-0.83). When compared to the random probability of selecting 599 the correct answer, the predictions of four classes are more correct than the predictions of two 600 classes. With two and four classes of predictions, the random probability is 1/2 and 1/4, respectively, so the accuracy above the random probability is higher for the four class 601 602 predictions than the two class predictions. The XGBoost models tend to perform slightly better 603 than the random forest models.



604

605 Figure 4

606 Cumulative accuracy of models when trained and tested on the same and different rock types. 607 To estimate the mean accuracy above the random probability (shown here with the text in the 608 boxes), we first calculate the accuracy above the random probability for the four machine 609 learning approaches (e.g., Figures S3, S4), and then we calculate the mean of this score from 610 the four approaches for the 36 combinations of training and testing rock types. The rows 611 indicate the rock type used to train the model. The columns indicate the rock type used to test 612 the model. The color is the mean accuracy above the random probability normalized by the 613 maximum value in each row. We expect to find high transfer accuracies for the 1) sandstone 614 and basalt, 2) monzonite and granite, 3) and limestone and shale. However, we only observe 615 this trend for 1) and 3). The other unexpected trend is the high transfer accuracies for monzonite, 616 shale and limestone.







619 Importance of groups of features that include information about the contraction (left), dilation 620 (center), and shear strain (right), for each rock type model and machine learning approach, for 621 features with importance >25% (a), >50% (b), and >75% (c) of the maximum importance of 622 each model. At the lowest threshold (a, 25%), most of the models depend on features with each 623 of the strain components, but the features with dilation tend to have the highest importance. At 624 the highest threshold (c, 75%), all of the models except the limestone models depend only on 625 the dilatation strain. For all of the other rock types and the machine learning approaches, 626 features with dilation have the highest predictive power of the strain components.





Importance of groups of features with each strain component (left) and statistic (right), averaged over all the rock types for each machine learning approach, for features with importance >25% (a), >50% (b), and >75% (c) of the maximum importance of each model. With increasingly important groups of features, dilation increases in average importance, while contraction and shear strain decrease in average importance. The intermediate values of the populations (25th-75th) tend to have the highest importance of the statistics.





635

636 Figure 7

637 Evolution of the magnitude of incremental strain components throughout three experiments: a) 638 sandstone experiment FBL01, b) monzonite experiment MONZ05, c) limestone experiment 639 ANS02, shown as the mean of the incremental strain population divided by the incremental 640 macroscopic axial strain (vertical axis), relative to the normalized macroscopic axial strain 641 (horizontal axis). The strain components identified as highly important in each rock type model 642 (dilation for sandstone and monzonite, contraction for limestone) tend to have higher mean 643 values than the other strain components. Only the dilation in the sandstone experiment (a) 644 shows a systematic evolution with increasing axial strain, suggesting that this component best 645 predicts the proximity to failure, as confirmed by the machine learning methods. In contrast, 646 the dilation in the monzonite (b) and contraction in the limestone (c) experiments do not show 647 a systematic evolution with distance toward failure, although the machine learning methods 648 identify these strain components as the best predictors of the proximity to failure.



650 **Figure 8**

651 Development of precursory inelastic dilatational strain toward catastrophic rupture. The 652 intermediate values of the dilatational strain population provide the best predictive power about 653 the proximity to macroscopic failure in triaxial compression experiments at the conditions of 654 the upper crust in 24 models with six rock types and four machine learning approaches.