

1 **Precise microbeam dating defines three Archaean granitoid suites at the southwestern
2 margin of the Kaapvaal Craton.**

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14 **ABSTRACT**

15 Precise microbeam U-Pb zircon dates have been determined for 17 granitoid samples which
16 crop out along the southwestern margin of the Kaapvaal Craton. In the Marydale High two
17 main types of granitoid are distinguished mainly by their normative Quartz - Alkali Feldspar –
18 Plagioclase mineral proportions. The Draghoender Granite type is generally tonalitic to
19 trondhjemite and some samples have very low heavy Rare Earth Elements, probably
20 originating as melts of eclogitic protoliths. Their ages vary by ~50 Ma from 2946 ± 9 Ma to
21 2892 ± 6 Ma (2σ), reflecting a long period of magmatism possibly due to subduction. The
22 granodioritic to monzogranitic Skalkseput Granite type intrudes the Draghoender granite in
23 places and has a coherent age of 2901 ± 14 Ma. These two granite types thus overlap in time
24 and space. The more evolved Skalkseput type could have been derived by melting of the
25 Draghoender type during assembly of the Kaapvaal Craton by collision of the Kimberley and
26 Witwatersrand Terranes.

27 A third granite type is exemplified by the 2721 ± 6 Ma monzogranitic Steenkop Granite
28 Gneiss which occurs south of Prieskapoort, and corresponds in age to a 2718 ± 8 Ma
29 monzogranite dyke which intrudes a 2871 Ma Skalkseput monzogranite. The previous age of
30 2718 ± 8 Ma for the Skalkseput Granite probably also dated a Steenkop-type dyke. A basal
31 conglomerate to the Zeekoebaard Formation of the Ventersdorp Supergroup has a single 2720
32 ± 4 Ma detrital zircon population.

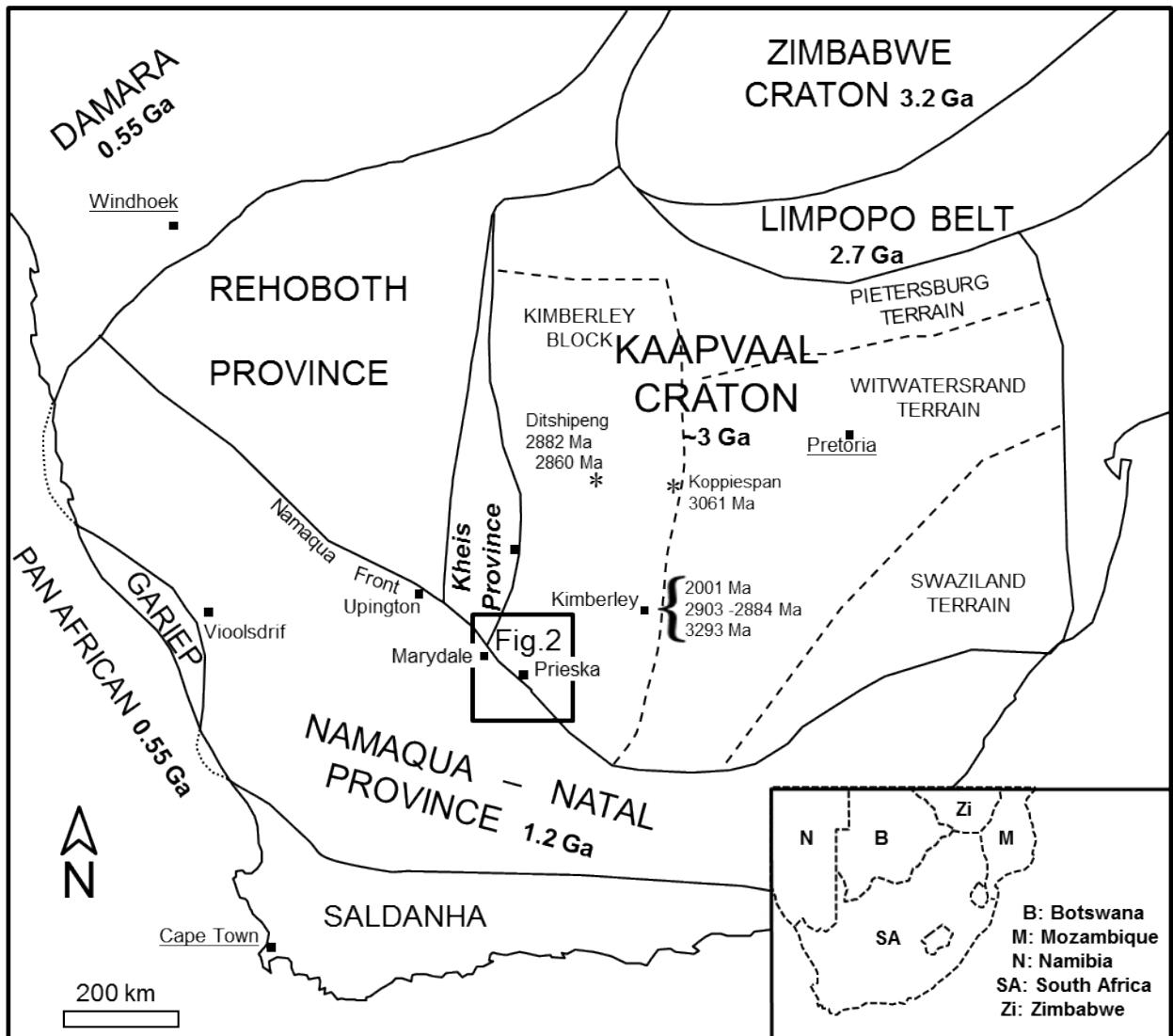
33 South of the Marydale High a narrow strip of Draghoender-type Granite extends to
34 Franzenhof where it is exposed in a Dwyka glacial outwash surface and is dated at 2931 ± 9

35 Ma. The southernmost exposures of granite are windows in the Dwyka group about 130 km
36 southeast of Marydale on both sides of the Doornberg Fault. The 2905 ± 7 Ma Maritzdam
37 Granite which is basement to Ventersdorp Supergroup volcanic rocks at Kuip is
38 monzogranitic and corresponds to the Skalksepuit type. On the southwest side of the
39 Doornberg Fault the 2907 ± 4 Ma Welgevonden Granite underlies Marydale Group volcanic
40 rocks, showing that the Kaapvaal basement extends across the fault.

41

42 Keywords: Kaapvaal basement; Archaean granite; U-Pb microbeam zircon dating; Lu-Hf
43 zircon;, crustal evolution; Kimberley Terrane.

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Figure 1. Precambrian structural provinces of southern Africa showing the locality of the study area (Fig. 2) at the southwestern margin of the Kaapvaal Craton. Selected precise U-Pb ages of other granitoids in the basement of the Kimberley Terrane are shown.

69

70 1. Introduction

71 The southwestern margin of the Kaapvaal Craton is marked by a dramatic transition from the
 72 undeformed Archaean basement granites and cover sequence of the craton to highly deformed
 73 and metamorphosed granitic and supracrustal rocks of the Mesoproterozoic Namaqua-Natal
 74 Province, shown in Fig. 1. The Marydale Terrane shown in Figure 2 lies between the two
 75 provinces and corresponds broadly to the dextral Doornberg Lineament of Vajner (1974b),
 76 bounded by the Doornberg Fault and Brakbos Shear Zone.

77 Granites are exposed along the Doornberg Fault. North of the fault they are unconformably
 78 overlain by the Kaapvaal volcanosedimentary cover sequence comprising ~2.7 Ga

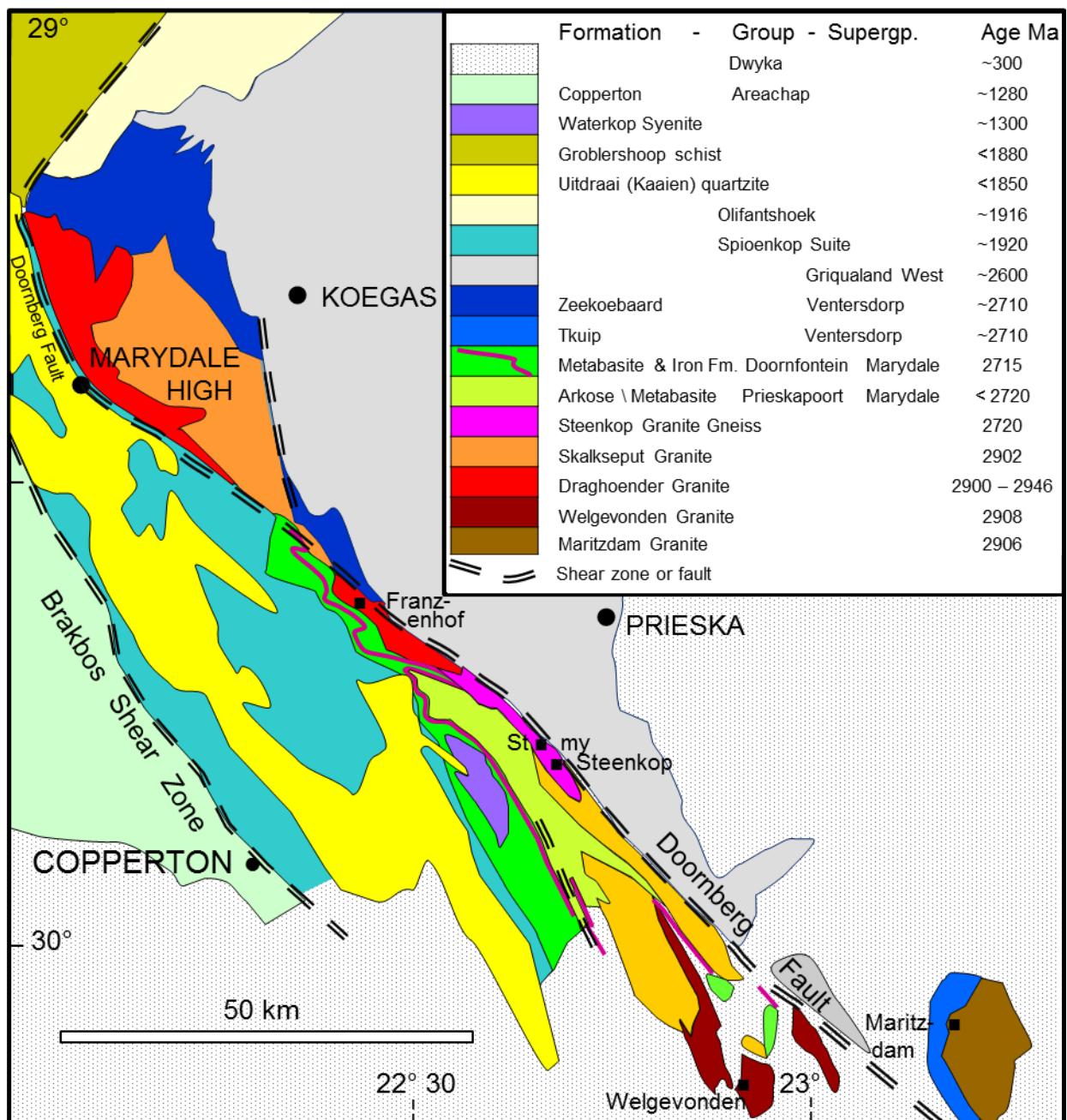


Figure 1. Geological map showing the granites exposed on the southwestern margin of the Kaapvaal Craton. The Marydale Terrane lies between the Doornberg Fault and Brakbos Shear Zone. The Marydale High is the area underlain by the Draghoender and Skalkseput Granites between Marydale and Koegas. Sample localities south of the Marydale High are indicated as follows. Franzenhof Granite DC1517; Steenkop Granite Gneiss DC00121; St my: Mylonitised Steenkop Granite DC01120 from Prieskapoort; Welgevonden Granite DC1333; Maritzdam Granite from the T'kuip window of basement granite overlain by Ventersdorp Supergroup volcanic rocks.

89 Ventersdorp and ~2.6 -2.2 Ga Griqualand West Supergroups and are clearly Kaapvaal
 90 basement granites. Granites also crop out south of the Doornberg fault in the Marydale
 91 Terrane, but their stratigraphic affinity is not obvious. In some places these granites are
 92 overlain by arkoses and metabasalts of the Marydale Group, but it has not been established
 93 whether the contact is unconformable or tectonic. The Marydale stratigraphy dips south-
 94 westwards but increases in grade upwards from greenschist to granulite, indicating that it is a
 95 thrust sequence, as is clearly seen in the isoclinal nappes of the Uitdraai Quartzites at the
 96 structural top of the sequence.

97 In this work we investigate the geochronology and geochemistry of the granites of this region
 98 in some detail and their relations across the Doornberg Fault. We establish the existence of
 99 three suites, each of which has characteristic composition and age.

100

101

102 ***1.1 The Marydale high***

103 By far the largest exposure of granites in the area, between Marydale and Koegas (Fig. 2), is
 104 termed the Marydale high (Stowe, 1986, Altermann and Halbich 1991), because it has clearly
 105 been uplifted relative to the Kaapvaal supracrustal rocks to the east. To the east it is
 106 overridden by thrusts of Spioenkop Suite metabasites and Uitdraai quartzites. Vajner (1974a)
 107 mapped and distinguished two granites in the Marydale High which represent the Kaapvaal
 108 basement because they are unconformably overlain by the Seekoebaard andesites of the ~2.7
 109 Ga Ventersdorp Supergroup. The supposedly older Draghoender Granite in the north-western
 110 part is rather heterogeneous, the normal type being a white to grey muscovite-biotite granite,
 111 commonly deformed and containing pegmatitic veins. The Skalkseput Granite which intrudes
 112 the Draghoender in places, is a more homogeneous light grey biotite muscovite granite and
 113 frequently K-feldspar porphyritic. The two varieties are not easy to distinguish in the field.
 114 The boundary between the two types is shown in Figure 4, following Malherbe and Moen
 115 (1991), which differs slightly from that of Vajner (1974b). Large xenoliths of Draghoender-
 116 like granite were found in the Skalkseput Granite by Vajner (1974a).

117

118 ***1.2 Previous ion probe dating in the Marydale High***

119 Three granite samples were dated by the U-Pb zircon ion probe method, reported in an
 120 abstract by McCourt et al., (2000), using data from a thesis by Hilliard (1999). They reported
 121 a ^{207}Pb - ^{206}Pb age of 2853 ± 4 Ma for a Draghoender Granite sample. A Skalkseput Granite
 122 sample gave a 3111-2930 Ma spread of ages and another sample previously analysed by R.

123 Armstrong gave an ^{207}Pb - ^{206}Pb age of 2718 ± 8 Ma (but no data are available for this sample).
124 Precise coordinates have not been documented for these samples, two of which are located as
125 sample numbers on the regional map of Hilliard (1999). He reported an intrusive contact
126 between what he termed the Skalksepuit Granite and the Zeekoebaart Formation (Ventersdorp
127 Supergroup) and ‘assuming that the Skalksepuit Granite comprises only one generation of
128 granitoid intrusives’, concluded that the Zeekoebaart Formation is older than 2718 ± 8 Ma.

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130 **2. Methods**

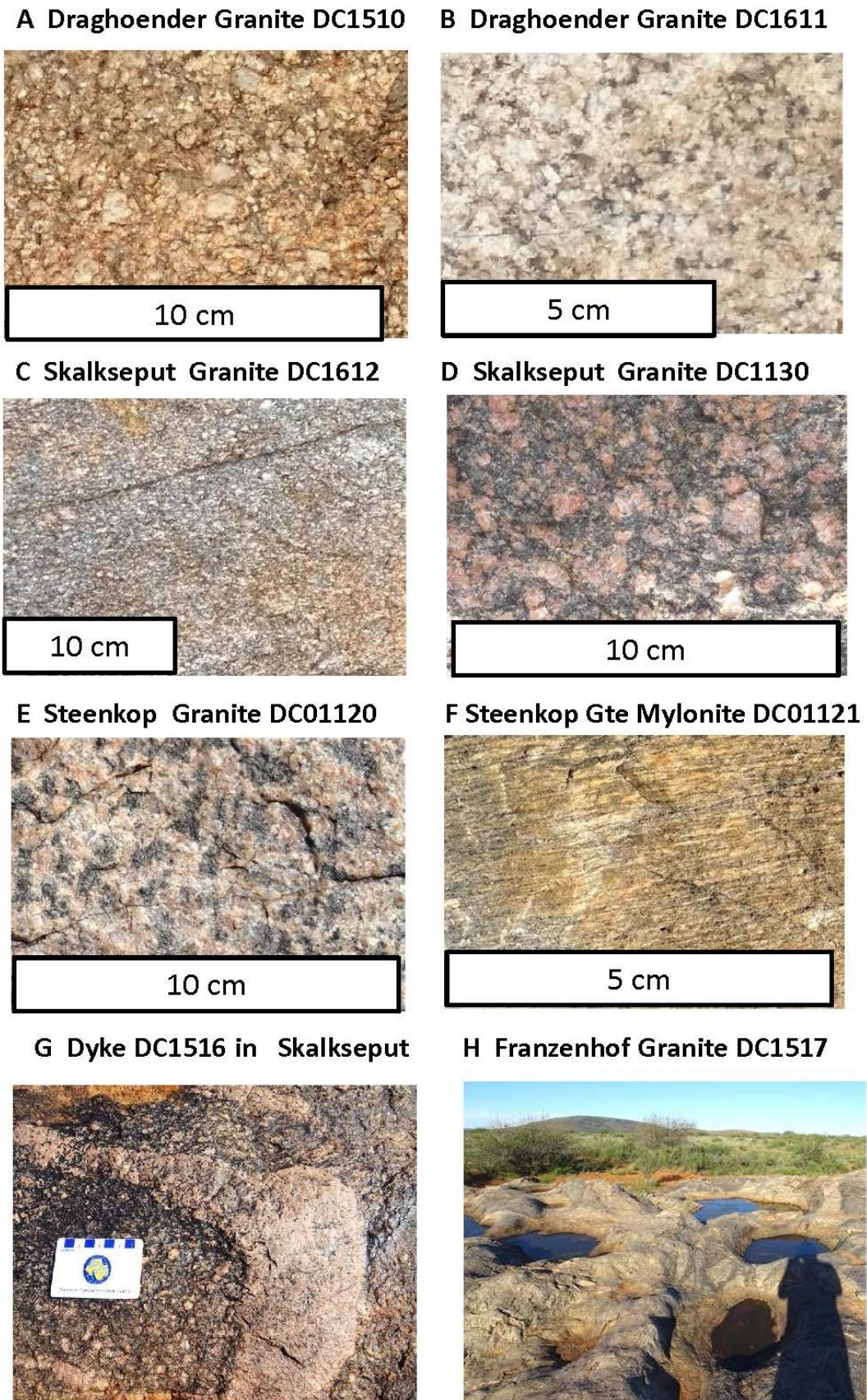
131 Samples were collected as about 1 kg of cm-sized fragments. They were coarsely crushed in a
132 chrome steel swing mill and a whole rock aliquot separated by cone-and halving, then finely
133 ground. Zircon was extracted from the rest of the samples by crushing to pass a 400 micron
134 sieve, then washed and manually panned to concentrate the heavy minerals. Zircon grains
135 were hand-picked and mounted in epoxy pucks for SEM imaging.

136 Geochemical analyses were done by commercial labs using XRF and ICPMS and given in
137 Table 1 and Appendix A1.

138 Microbeam U-Pb zircon dating was done at three labs by methods briefly described below. For
139 more detail see Cornell et al, (2016). The complete U-Pb data are given in appendix A2. At the
140 University of Oslo, Norway Laser ablation multicollector ICPMS zircon dating was done on
141 three samples using a 213 nm Nd-YAG laser ablation sampler with 40 micron spot size and
142 analysed for U and Pb isotopes, following the method described by Andersen et al., (2009).
143 Common lead corrections were not applied due to the difficulty of distinguishing ^{204}Pb from
144 ^{204}Hg in the system, but points with high common lead were omitted, identified by low
145 $^{206}\text{Pb}/^{204}\text{Pb}$ ratios and ^{204}Pb signal higher than the ^{204}Pb -free zircon standard 91500. The same
146 zircon grains were analysed for Lu-Hf with a 60 micron rectangular spot as described by
147 Cornell et al. (2012), and the data is presented in the appendix A3.

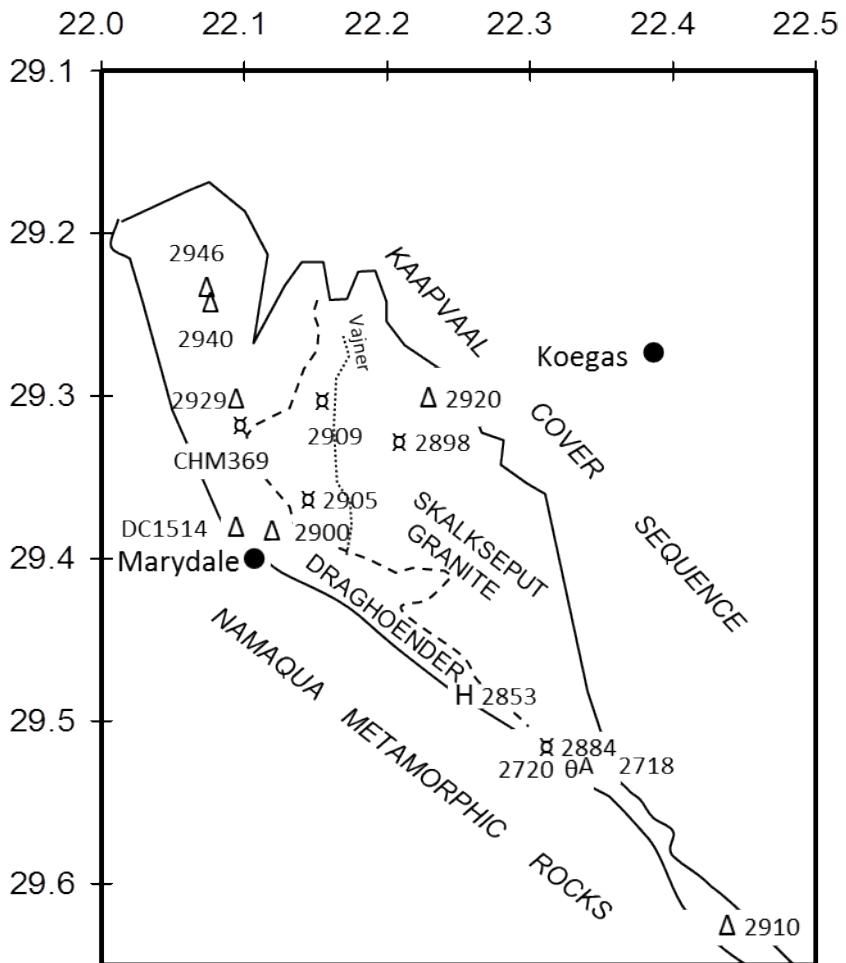
148 Laser Ablation Inductively Coupled Plasma high-resolution mass spectrometric data for 11
149 samples were obtained at the Central Analytical Facility, Stellenbosch University, employing
150 a Thermo Finnigan Element 2 mass spectrometer coupled to an ASI Resolution M50-E
151 excimer laser ablation system. Single spot analyses had 20 μm diameter and crater depth 10-
152 15 μm . The methods employed for analysis and data processing are similar to those described
153 by Frei and Gerdes (2009) and common lead corrections were made where required.

154



155

156 Figure 3. Field pictures of samples taken in this work. Image H shows a rejuvenated Dwyka
 157 glacial outwash surface at Franzenhof.



158

159 Figure 4. Precise U-Pb zircon dates for Kaapvaal basement granitoid samples exposed
 160 in the Marydale High, with localities marked by symbols. Δ indicates Draghoender type
 161 tonalitic rocks with low QAP normative alkali feldspar (4-12%). \square indicates Skalkseput –
 162 type granodioritic rocks with higher QAP normative alkali feldspar (22-29%). θ indicates
 163 younger (2720 Ma) monzogranite dykes coeval with the Steenkop Granit. The sample
 164 numbers corresponding to ages can be found in Table 2. Samples for which norms but not
 165 dates are available are shown with sample numbers. The H 2853 Ma age shown is by ion
 166 probe for Draghoender Granite from Hilliard 1999 and the A 2718 Ma age shown is an ion
 167 probe age by Richard Armstrong for his Skalksepets Granite, referred to in Hilliard 1999 and
 168 McCourt et al., (2000) and considered in this work to represent a Steenkop – aged dyke
 169 which we sampled at this locality. Precise locations for these two samples are not available.
 170 The dashed line is the Draghoender – Skalkseput boundary according to the Council for
 171 Geoscience 1:250 000 Prieska map sheet of Malherbe and Moen 1991. Where it differs from
 172 that, the boundary according to Vajner 1974a is shown as a dotted line.

173

174

175 Th-U-Pb zircon dating was also done for three samples using the NordSIM ion probe as
176 described by Whitehouse and Kamber (2005). A ~10 micron spot size was used and the NIST
177 91500 zircon standard was used for calibration. Common Pb corrections assume a present day
178 Stacey and Kramers (1975) model average terrestrial Pb composition. The ion probe U-Pb
179 data are given in the appendix A2. All U-Pb age calculations were made using the Isoplot 3
180 programme of Ludwig (2012). Uncertainties of age calculations are all given at the 2σ level,
181 ignoring decay constant errors.

182

183 **3. Samples in the Marydale High**

184 Twelve granite samples were taken from the Marydale High. Their localities are shown in
185 Figure 3. Eleven of them were originally classified as Draghoender or Skalkseput Granite
186 according to the map of Malherbe et al, (1991), and their stratigraphic affinity is further
187 examined below. The sample were generally taken from small outcrops which were separated
188 by large areas of sub-outcrop. Some of the outcrops showed morphological features such as
189 polished surfaces and potholes suggesting that they were Dwyka glacial outwash surfaces.
190 One sample DC1516 was taken from a small group of meter-wide dykes which intrude
191 Draghoender Granite near a prominent road cutting in the south of the Marydale High and this
192 is thought to be the rock which McCourt et al. (2000) regarded as Skalkseput Granite. Near
193 this locality (110 m northeast) is a small outcrop of deformed conglomerate comprising
194 quartzite pebbles in a micaceous matrix. The field relationships with the surrounding granites
195 were not apparent and this could represent either a xenolith or an unconformably overlying
196 part of the Zeekoebaard Formation, regarded as a Ventersdorp Supergroup correlate. This
197 was also sampled for detrital zircon dating (DC1515).

198

199 **4. Granites southward astride the Doornberg Fault**

200 The Marydale high is bounded on both sides by branches of the Doornberg Fault, which dies
201 out northward on the east side but continues past Marydale on the west (Figure 1), where
202 intense late movements on the fault generated thin veins of pseudotachylite in the Marydale
203 amphibolites (Cornell, 1974). South of the Marydale high the main Doornberg Fault is a
204 single shear zone in which several phases of deformation can be recognized e.g. at
205 Prieskapoort (Scott, 1987). Granitoids are exposed on both sides of the fault. An important
206 question is whether those to the southwest are the same as those to the northeast. To
207 investigate this, samples were taken at four localities southwest of the fault, indicated by
208 letters in Figure 1. At Franzenhof (Fr) sample DC1517 was taken from a Dwyka glacial

209 outwash surface with polished rounded topography and many potholes as showjn in Figure
210 3H.

Sample equivalent	DC1510	DC1512	DC1513	DC1514	DC1516	DC1517	DC1518	DC1610	DC1611	DC1612	DC1613	CHM365	CHM369	CHM371	CHM372
Sample taken as	Drøghoender Granite	Skalkseput Granite	Skalkseput Granite large outcrop	Drøghoender Granite	Dyke in Skalkseput Granite	Franzenhof fluvioglacial surface	Marizdam Granite at Kuijip	Drøghoender Granite	Drøghoender Granite	Skalkseput Granite	DC1130				
Geochem															
Type	Drøghoender	Skalkseput	Skalkseput	Drøghoender	Stenktop	Drøghoender	Stenktop	Drøghoender	Drøghoender	Skalkseput	Drøghoender	Stenktop	Skalkseput	Skalkseput	Skalkseput
Lat. °S	29.2444	29.3643	29.3042	29.3860	29.5159	29.6283	30.1425	29.2364	29.3041	29.3288	29.8110	29.31759	29.3857	29.5163	29.31109
Lon. °E	22.0746	22.1443	22.1529	22.0853	22.3112	22.4341	23.2251	22.0716	22.0904	22.2247	22.6850	22.1171	22.09434	22.3109	22.31109
SiO ₂	70.11	74.14	74.24	72.36	75.93	73.05	70.08	70.67	69.05	75.24	75.93	74.49	73.14	71.00	77.91
TiO ₂	0.24	0.14	0.11	0.25	0.12	0.21	0.45	0.24	0.29	0.05	0.08	0.58	0.08	0.16	0.10
Al ₂ O ₃	16.29	14.27	14.47	15.44	12.57	14.92	14.66	15.75	16.02	13.86	13.77	12.37	14.16	14.24	14.53
Fe ₂ O ₃	2.05	1.68	1.28	2.20	1.73	1.73	3.04	1.99	2.04	0.51	0.67	3.17	1.36	2.31	0.67
MnO	0.035	0.027	0.033	0.028	0.020	0.042	0.040	0.040	0.020	0.010	0.020	0.045	0.021	0.034	0.01
Mo	0.57	0.14	0.03	0.32	<0.01	0.40	0.67	0.040	0.02	0.01	0.02	0.30	0.01	0.38	0.33
CaO	3.06	1.17	0.78	2.77	0.39	1.82	1.34	0.58	0.72	0.07	0.24	1.40	0.81	2.31	0.62
Na ₂ O	5.42	4.41	4.70	5.56	5.02	3.87	2.92	3.51	4.41	1.43	4.41	4.15	5.12	4.01	4.48
K ₂ O	0.99	3.60	3.70	0.72	5.24	1.84	4.88	5.29	5.66	4.64	5.49	4.15	4.49	4.54	3.77
P ₂ O ₅	0.094	0.054	0.057	0.111	0.036	0.120	0.227	0.980	0.950	4.070	1.060	0.029	0.029	0.144	0.050
Cr ₂ O ₃	0.007	0.005	0.006	0.004	0.005	0.007	0.005	0.090	0.110	0.100	0.020	0.007	0.005	0.015	0.010
Li ₂ O	0.87	0.33	0.42	0.58	0.30	0.88	0.71	1.08	1.04	0.55	0.67	0.29	0.57	0.47	0.68
Total	99.97	99.82	100.35	99.56	100.05	99.97	99.63	99.41	99.51	99.38	100.06	100.11	99.69	100.03	100.84
CPW - normative minerals wt%															
Quartz	31.51	30.84	30.33	35.12	32.16	24.58	28.8	24.04	32.11	36.8	34.41	30.13	30.9	26.26	35.54
Plagioclase	61.1	42.96	43.55	60.29	30.24	51.25	38.33	59.67	64.52	41.07	34.95	39.11	54.79	40.63	43.9
Orthoclase	5.91	21.59	21.98	4.25	31.2	10.99	29.13	5.91	5.73	24.29	6.32	27.78	8.92	27.01	22.22
Conundum	0.98	1.13	1.47	0.76	1.65	1.11	0.89	0	1.34	1.07	0.08	0.86	0.99	0.37	0.31
Dolomite	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.3	0
Hypsthene	3.77	2.12	1.42	3	1.85	2.77	3.5	3.27	0.71	1.29	3.56	1.51	3.15	3.32	2.21
Ilmenite	0.46	0.28	0.21	0.47	0.23	0.4	0.85	0.46	0.57	0.09	0.15	1.1	0.15	0.4	0.19
Narneite	0.49	0.38	0.64	0.51	0.51	0.88	0.58	0.61	0.14	0.20	0.93	0.39	0.62	0.67	0.19
Apatite	0.23	0.12	0.14	0.25	0.09	0.28	0.53	0.21	0.25	0.05	0.35	0.07	0.23	0.32	0.19
rock type															
Tonalite	33	32	32	36	34	27	31	25	33	38	37	31	33	28	36
% in QAP	29	22	23	4	32	12	32	6	6	25	7	26	9	19	23
A ₉₀	6	22	23	45	64	31	54	42	45	37	40	58	43	53	41
P ₉₀	65	45													

Table 1. Major element analyses and normative data for granitoid samples from the Southwestern margin of the Kaapvaal Craton.

Stratigraphic unit taken	Lat. S	Lon. E	Deg	Zircon age domains seen in SEM images	method	lower	intercept	proba	xenocryst ages Ma	comments
						Age Ma	$\pm 2\sigma$	age Ma		
DC1610	Draghoender tonalite	29.236	22.072	Main Cl-grey with oscillatory zonation. A few cores CL-bright and truncated zonation.	discordia LA HR ICPMS	2946	9	13	120	1-3 cores analysed not Concordia age 2942 ± 8 agrees older
DC1510	Draghoender tonalite	29.244	22.075	Main Cl-grey with zonation. Cores CL-bright and truncated zonation.	discordia LA HR ICPMS	2940	6	66	77	1-3160-3000 Concordia age 2937 ± 7 agrees
DC1611	Draghoender tonalite	29.304	22.090	Mains Cl-dark, some zoned xenocryst cores	discordia LA HR ICPMS	2929	11	-18	140	0.42 3125 - 3020 Concordia age agrees: 2936 ± 15 Ma
DC1513	Skalkseput	29.304	22.153	Mains Cl-dark but oscillatory zoned, cores CL-bright	discordia LA HR ICPMS	2909	9	-1	12	1 none older Cores all coeval with mains, perfect recent lead loss model, one discordant point.
DC1613	Skalkseput	29.305	22.225	Mains Cl-dark, some cores seen in BS images	discordia LA HR ICPMS	2920	18	274	420	0.49 3390 - 3000 >15% discordant points show complex lead loss.
DC1612	Skalkseput	29.329	22.208	Mains Cl-dark, cores CL-bright	discordia LA HR ICPMS	2898	7	-13	53	1 none older Cores all coeval with mains, concordia age agrees: 2897 ± 7 Ma.
DC1512	Skalkseput	29.364	22.144	Mains Cl-grey, cores CL-bright	discordia LA HR ICPMS	2905	8	54	70	0.79 3155 & 3120 Cores mostly coeval with mains, concordia age agrees: 2903 ± 9 Ma.
DC1332	Draghoender hoender	29.386	22.117	Mains Cl-dark but oscillatory zoned, cores CL-bright & truncated, thin rims CL-bright	discordia SIMS	2900	6	-42	72	0.78 3040 Only 6 points of 15 conform, others show scatter probably reflecting complex lead loss. Overall regression: 2893 ± 17 Ma.
DC1515	Seekoebaard conglomerate	Steen age	29.515	22.312 Two main groups, CL-dark and CL-bright, generally unzoned.	discordia LA HR ICPMS <20% discordant	2720	4	97	95	1 none younger 2633 ± 34 Ma point >10% discordant Concordia agrees, remarkable single detrital age population suggests derived from Ventersdorp aged magmatic rocks.
DC01130	Skalkseput	Skalkseput	29.516	22.311 Mains Cl-dark & oscillatory zoned, cores CL-bright.	discordia LA MC ICPMS	2884	22	277	49	0.16 not analysed Ancient lead loss, 277 Ma corresponds to Karoo Dwyka glaciation and diogenesis.
DC1516	Dyke in Skalkseput (DC130)	Steen-kop	29.516	22.311 Mains Cl-dark with some zonation, few CL-bright xenocrysts.	discordia LA MC ICPMS	2720	10	287	200	0.997 one core same age as mains Ancient lead loss, more discordant points trend towards origin. Concordia age agrees: 2720 ± 10 Ma.
DC1517	Franzenhof granodiorite	Drag-hoender	29.628	22.434 Mains are CL-grey with slight zonation, cores are CL-bright with truncated zonation.	discordia LA HR ICPMS <20% discordant	2912	10	-72	8	0.69 3105-2945 Concordia age agrees; including >20% discordant points gives intercepts, 200 & 2930 Ma.
DC00121	Steen-kop mylonite	Steen-kop	29.785	22.667 Broken grains, both CL-dark and zoned, thin CL-bright or dark metamorphic rims	Pb-Pb age LA MC ICPMS	2720	10 fixed at 0	50	0.97 2859 - 2792 one grain 2126	Unconstrained lower intercept -166 ± 490 Ma. Concordia age agrees: 2715 ± 14 Ma.
DC01120	Steen-kop type locality	Steen-kop	29.810	22.688 Mains Cl-dark but oscillatory zoned, cores CL-bright with main rims	discordia LA MC ICPMS	2719	4 fixed at 0	50	0.62 2830	Unconstrained lower intercept -77 ± 320 Ma. Concordia age agrees 2719 ± 6 Ma. Agrees with SIMS analysis below.
DC1518	Ongers River	Skalk-seput	30.143	23.225 Mains Cl-dark with some metamict zones, a few discordant CL-grey cores	discordia SIMS	2720	5 fixed at 0	50	0.63 not analysed Unconstrained lower intercept 151 ± 180 Ma. Concordia age agrees 2722 ± 7 Ma. Agrees with LA-MC-ICPMS above.	
DC1333	Welgevonden	Skalk-seput?	30.158	22.908 Mains Cl-grey & zoned, some metamict zones; a few discordia SIMS CL-bright resorbed cores.	concordia LA HR ICPMS	2906	8	300	350	0.9 none older Ancient lead loss, >10% more discordant points scatter. Concordia age agrees: 2905 ± 7 Ma.
						2908	4	115	67	0.18 3051 & 3001 Concordia age agrees: 2907 ± 5 Ma.

213 Table 2. Summary of age determinations and calculations.

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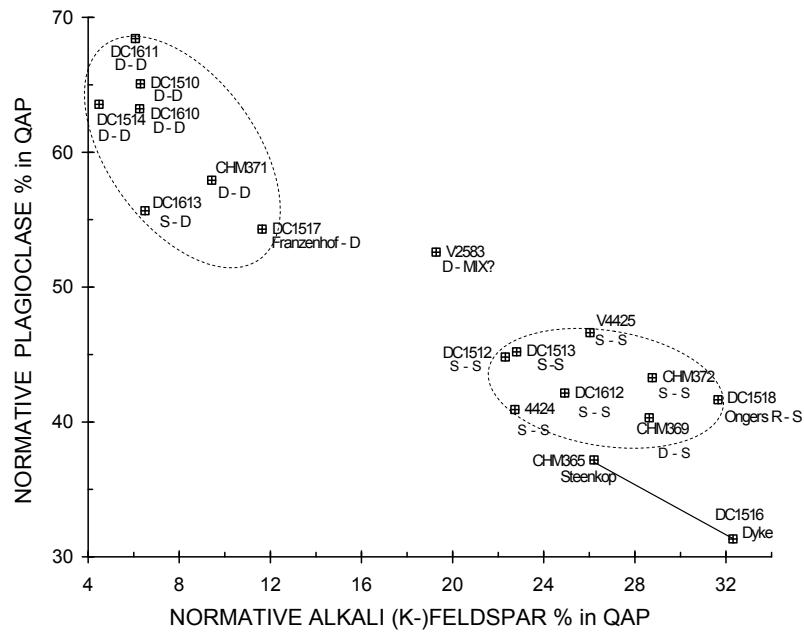
215 At Prieskapoort a mylonised variety of the Steenkop Granite was taken, sample DC00121 (St
216 my). At Steenkop (St) sample DC01120 was taken from the type locality of this biotite-
217 flecked monzogranite gneiss. A poorly exposed granite sample was taken on the farm
218 Welgevonden (sample DC1333, Wel), the southernmost exposure before Dwyka cover
219 becomes total. One sample from northeast of the Doornberg Fault was taken at the
220 Maritzdam, (sample DC1518), which underlies the exposure of Ventersdorp lavas at Kuip,
221 also an inlier in the Dwyka cover.

222

223 **5. Results**

224 CIPW norms for the samples analysed are given in Table 3 of the ancillary data. Data for
225 three granite samples from the Marydale High taken by Vajner (1974a) were also considered.
226 Diagrams based on normative mineralogy proved useful in distinguishing three granite types
227 as shown in Figure 5. Further geochemical discrimination using major elements is shown in
228 Figure 6 and a tectonic discrimination diagram in Figure 7. The U-Pb zircon microbeam data
229 are shown in concordia diagrams with age calculations in Figure 8. The zircon age domains,
230 details and comments about the geochronological results are summarized in Table 2. Rare
231 Earth Element diagrams are shown in Figure 9. The zircon Lu-Hf isotope data for twelve
232 samples are plotted in Figure 10. The CHUR values are from Bouvier et al. (2008) and the
233 DM values are from Griffin et al. (2004) modified to the CHUR values from Bouvier et al.
234 (2008) and the decay constant of ^{176}Lu from Söderlund et al. (2004). The geochemical and age
235 data are discussed together in the following section.

236



237

238 Figure 5. Normative Quartz – Alkali feldspar – Plagioclase (QAP) mineral proportions for
 239 Granite samples from the southwestern margin of the Kaapvaal Craton including the
 240 Marydale high and outcrops southwards on both sides of the Doornberg Fault. Three types are
 241 distinguished as shown. The Draghoender and Skalksepuit types overlap in U-Pb zircon age
 242 and the Steenkop type samples are significantly younger at ~2720 Ma.

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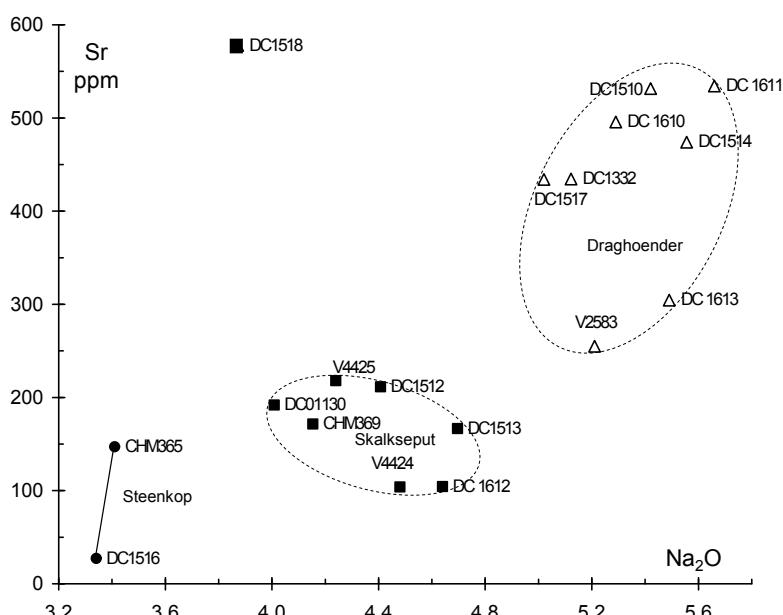
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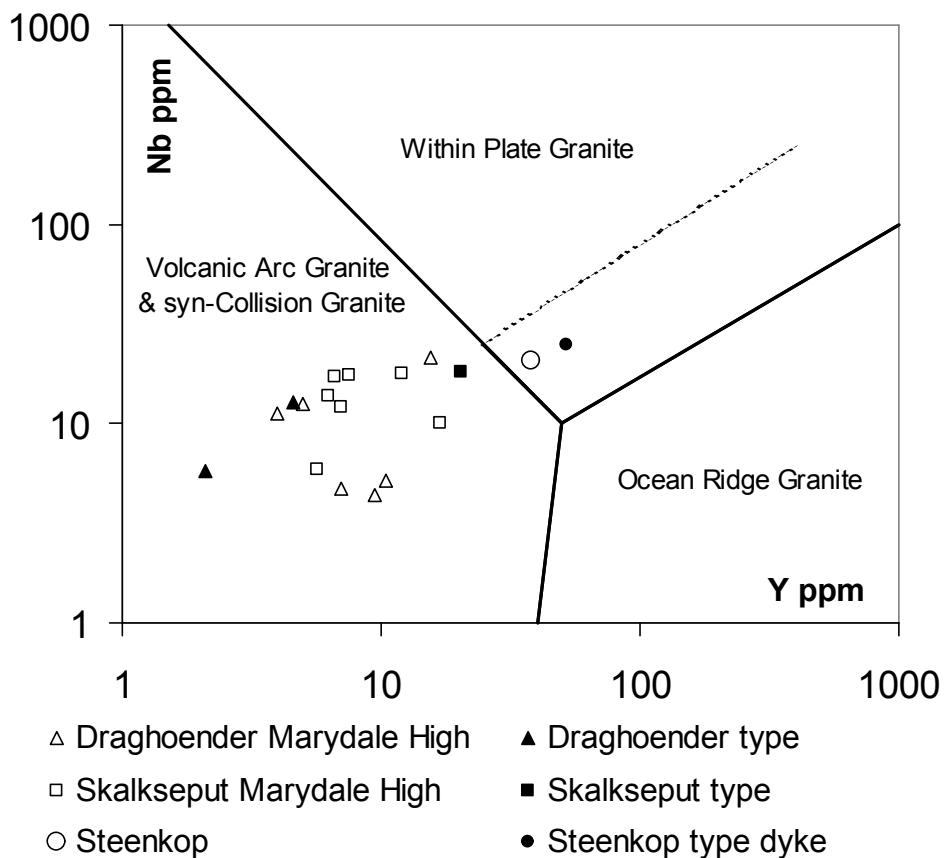
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259

260 Figure 6. Geochemical distinction of granite samples from the Southwestern Kaapvaal Craton.
 261 Three types are apparent, the Draghoender (triangles) and Skalksepuit (blocks) granites which
 overlap in age and the younger ~2720 Ma Steenkop type.



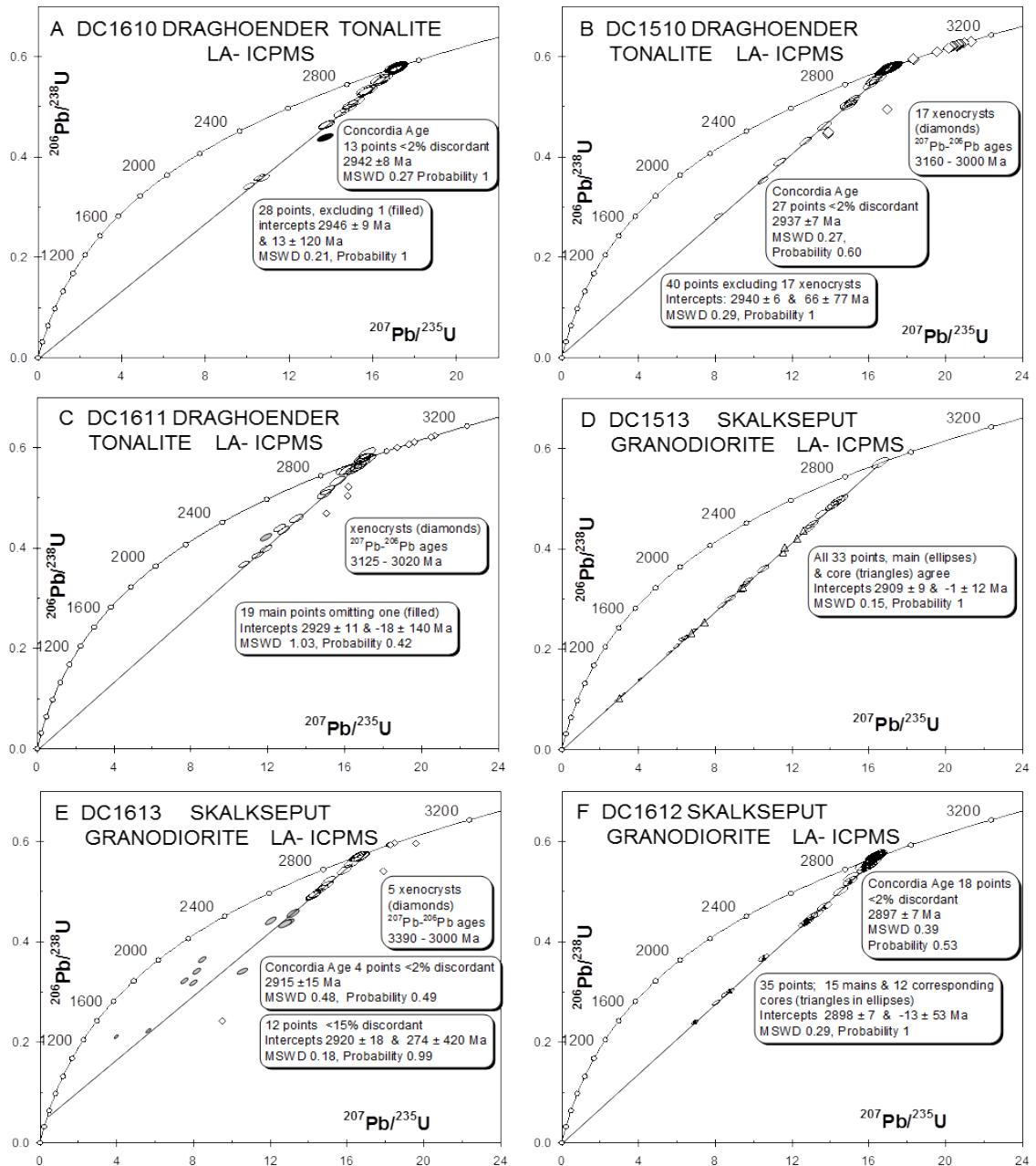
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263 Figure 7 Tectonic discrimination diagram for granites.

264

265 **6. Discussion**266 ***6.1 Stratigraphic identity of Draghoender and Skalkseput Granites***

267 There is some uncertainty about the boundary between the Draghoender and Skalkseput
 268 Granites in the Marydale High. Vajner (1974a) and Malherbe et al, (1991), mapped different
 269 boundaries in the central Marydale High as shown in Figure 4. We used the latter boundary
 270 which corresponds best to our results. The stratigraphic affinity of samples taken further south
 271 on either side of the Doornberg Fault is also not clear. Criteria were sought which could
 272 distinguish the Draghoender and Skalkseput Granites. Normative mineralogy proved to be
 273 useful, in that most of the samples taken as Draghoender Granite classify as tonalites on a
 274 normative Quartz-Akali Feldspar-Plagioclase (QAP) plot, whereas most of the Skalkseput
 275 samples are granodiorite or monzogranite with 22 to 29% A. This is clearly shown in the A-P
 276 plot in Figure 5, which is labelled with sample numbers and symbols e.g. D – S indicating
 277 taken as Draghoender, classified as Skalkseput type, with 4.5 to 13% A,
 278



280

281

Fig. 8-1. Concordia diagrams showing the U-Pb zircon data and calculations for six samples from the Marydale High, arranged in order of latitude south, following Table 2. Error ellipses are all shown at 1σ but age calculations at the 2σ confidence level.

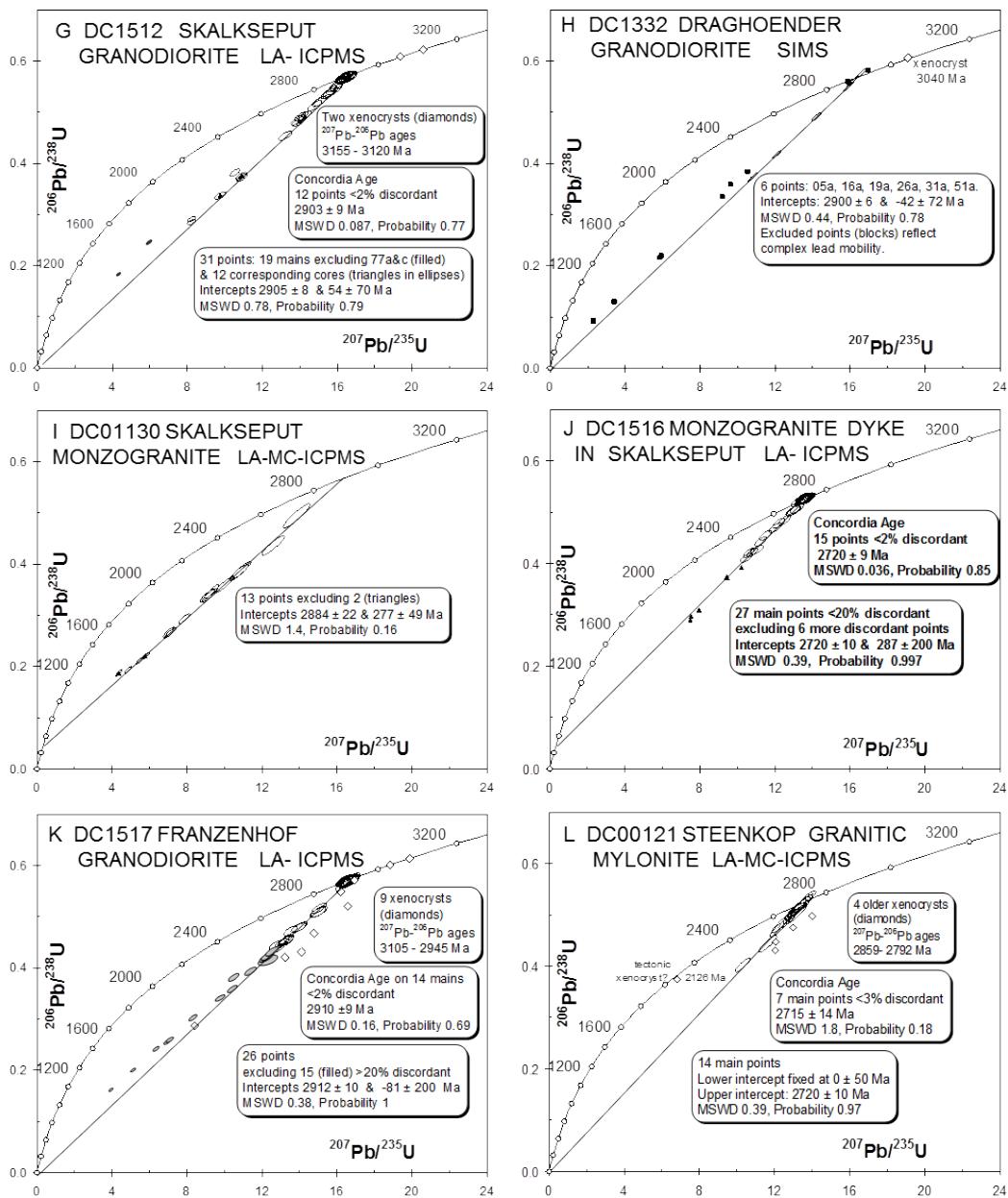


Fig. 8-2. Concordia diagrams showing the U-Pb zircon data and calculations for four samples from the Marydale High and two samples from further south, following Table 2. Error ellipses are all shown at 1σ but age calculations at the 2σ confidence level.

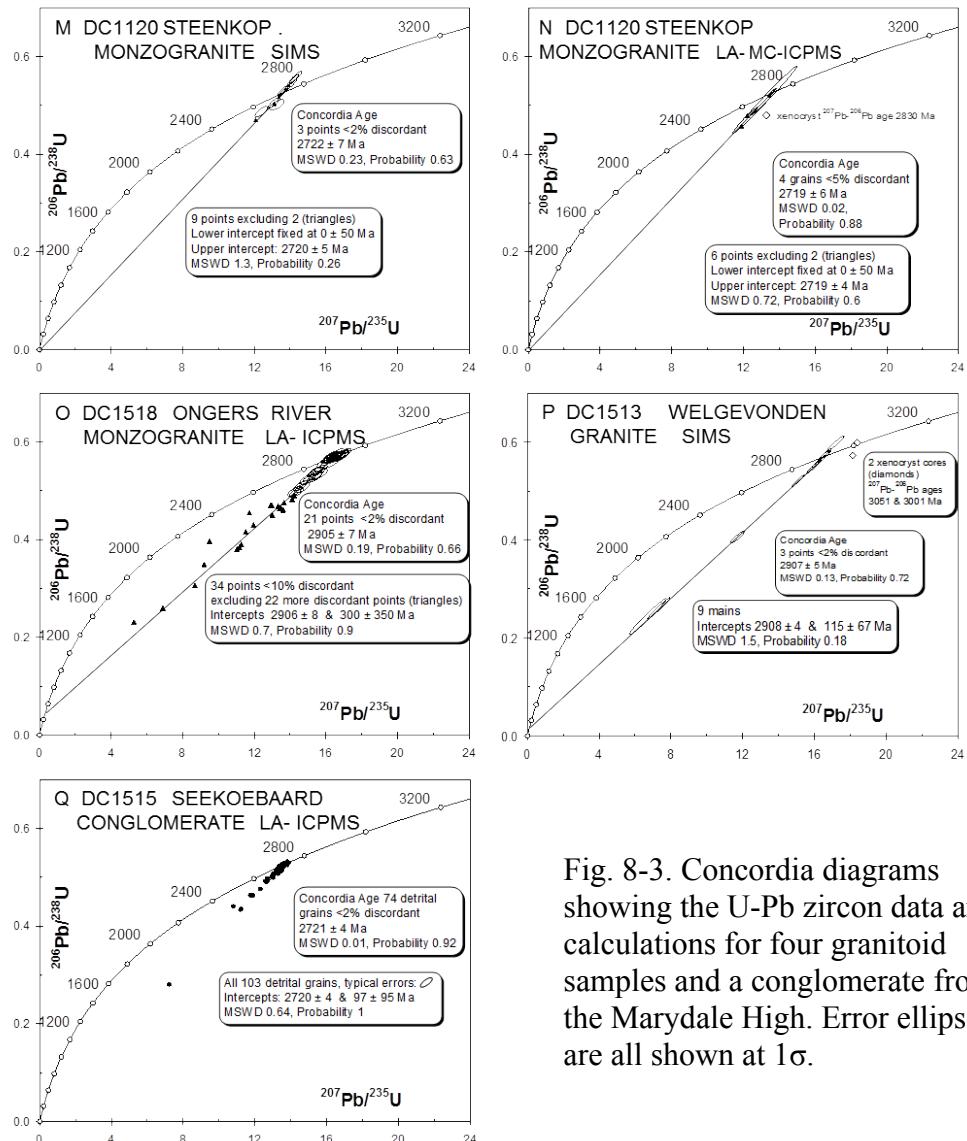


Fig. 8-3. Concordia diagrams showing the U-Pb zircon data and calculations for four granitoid samples and a conglomerate from the Marydale High. Error ellipses are all shown at 1σ .

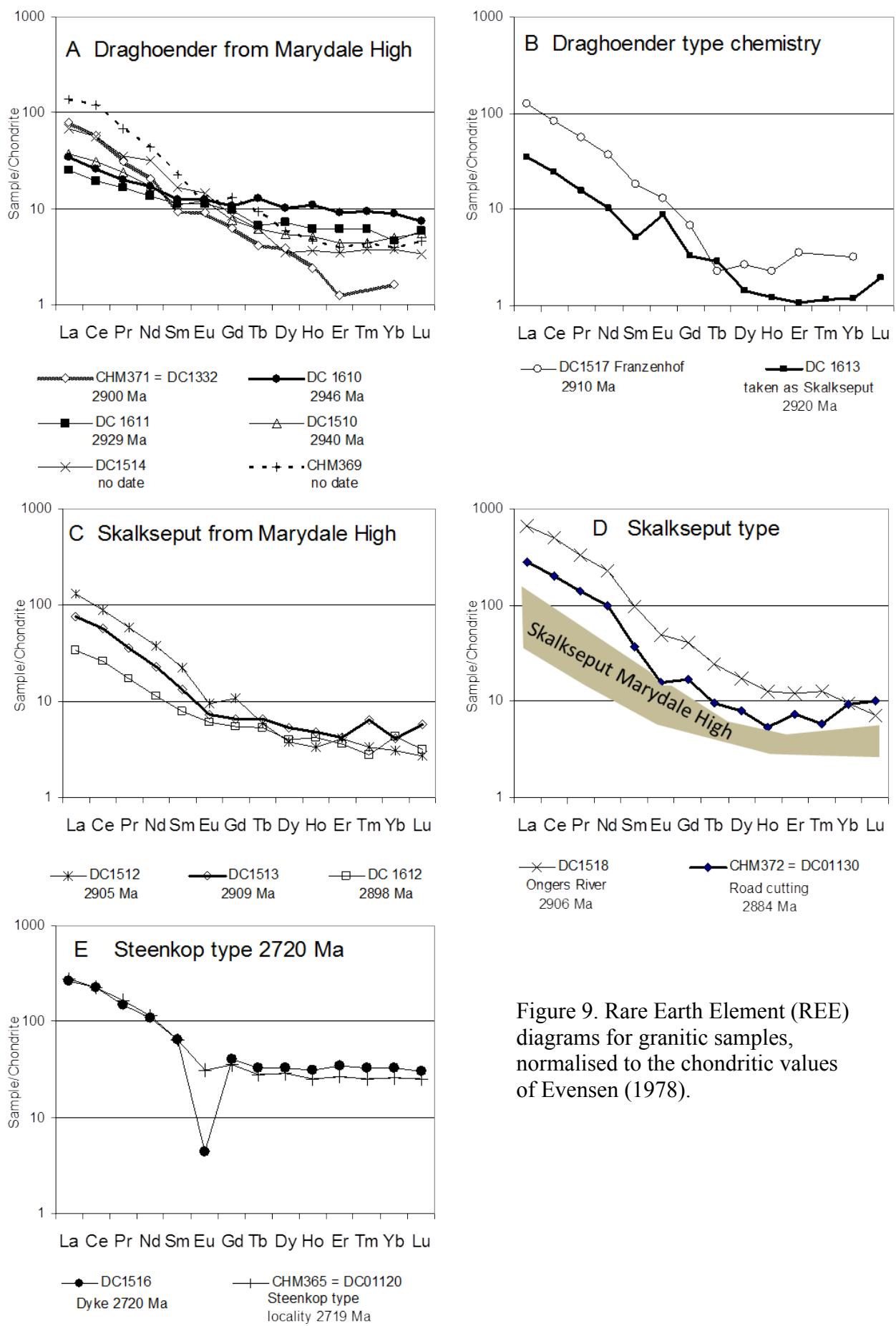
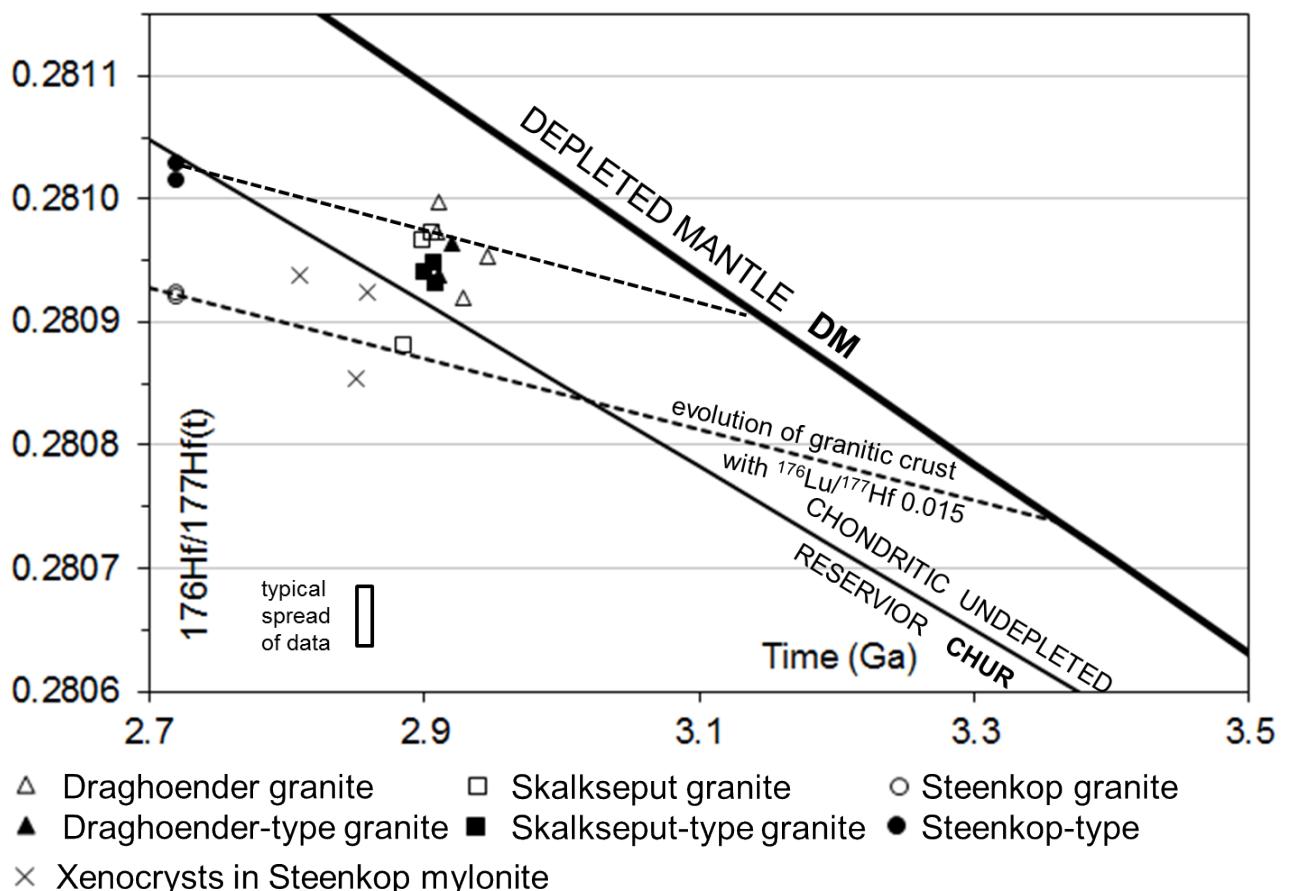
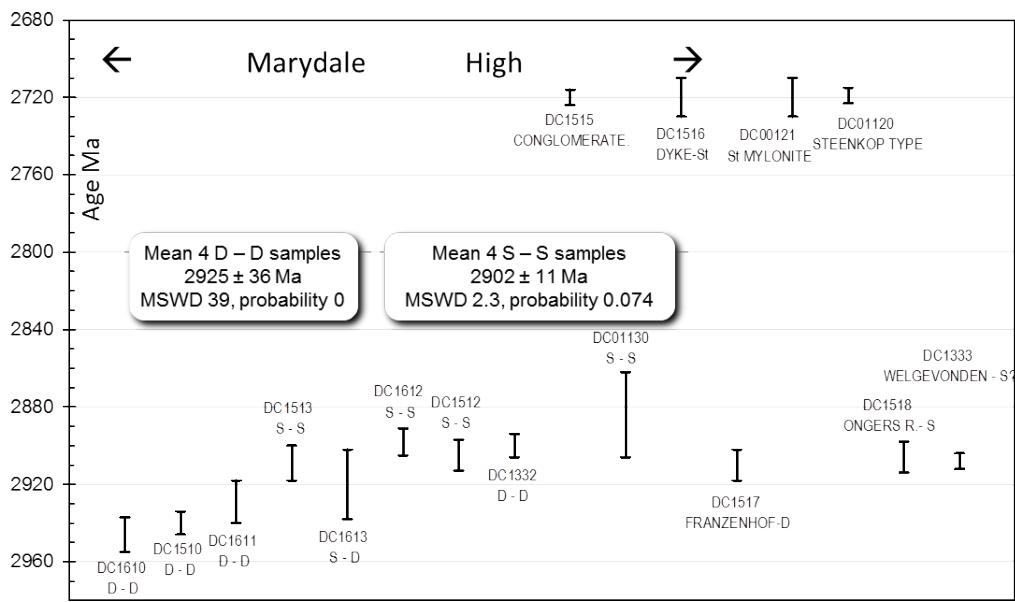


Figure 9. Rare Earth Element (REE) diagrams for granitic samples, normalised to the chondritic values of Evensen (1978).



286
287 Figure 10. Zircon Lu-Hf data for 12 samples plotted on a Hf isotope evolution diagram. The
288 data points are averages based on at least four Lu-Hf isotope analyses with typical spread \pm
289 0.000035. The calculation of crustal residence ages for a sample is illustrated by projecting an
290 evolution line corresponding to a hypothetical granitic precursor with $^{176}\text{Lu}/^{177}\text{Hf} 0.015$ from
291 the data point back to intersect with the Depleted Mantle line, in this case at 3.5 Ga. More mafic
292 precursors with e.g. $^{176}\text{Lu}/^{177}\text{Hf} 0.022$ would give older crustal residence ages.
293

294
295 There are clearly two groups (excluding for the moment the younger Steenkop type) and only
296 Vajner's (1974a) Draghoender sample plots between them. According to this grouping, one
297 sample DC1613, taken as Skalksepups plots in the Draghoender group, whereas CHM369
298 taken as Draghoender, plots in the Skalkseput group. To test whether the geochemical
299 grouping is supported by the geochronology, the ages of all the samples are plotted on a map
300 in Figure 4, and shown in order of latitude in Figure 11, in which e.g. S-D indicates taken as
301 Skalkseput, grouped as Draghoender.



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Figure 11. Precise U-Pb zircon microbeam dates for granitic samples from the southwest margin of the Kaapvaal Craton arranged in order of latitude south. Sample numbers are followed by Taken As - Classified As. e.g D – S means taken as Draghoender, classified as Skalksepuit type. St means Steenkop type.

308
309 The four Draghoender samples about which there is no doubt have ages between 2946 and
310 2900 Ma. They do not give a valid mean age and there seems to be a trend of decreasing age
311 southwards. The easterly sample DC1613 with age 2920 ± 18 Ma, which was taken as
312 Skalksepuit Granite but plots as Draghoender type, might represent a large xenolith or
313 intrusive body different from the Skalksepuit Granite; however its field relations could not be
314 determined in this area of poor outcrop. Tonalitic rocks of the Draghoender Granite type
315 clearly intruded over a period of about 40 Ma and were not comagmatic but probably
316 originated by similar magmatic processes.
317

318 Four unequivocally Skalksepuit samples give a mean age of 2902 ± 11 Ma with low
319 probability which could be valid if the errors were increased. Sample CHM369, taken as
320 Draghoender but plotting geochemically with Skalksepuit, was not dated. To summarise, the
321 emplacement of the Draghoender Granite type took place over a long period between 2946
322 and 2900 Ma. The Skalksepuit Granite might represent a single magmatic event at 2902 Ma,

323 which overlapped in time with the emplacement of the younger members of the Draghoender
 324 type.

325

326 ***6.2 The Steenkop type***

327 The Steenkop Granite Gneiss is about 200 Ma younger than the other granites with a weighted
 328 mean age of 2719 ± 3 Ma on three analyses of two samples including a mylonised sample
 329 DC00121 from the Doornberg Fault. The distinctive gneiss with large biotite flecks has a
 330 rather restricted outcrop, but the dyke (sample DC1516) which intrudes Skalksepuit Granite in
 331 the Marydale High has an indistinguishable age of 2720 ± 10 Ma, although it is a much lighter
 332 colour. This is also the age reported by McCourt et al., (2000) for the Skalksepuit Granite
 333 dated by Richard Armstrong. Failing any input from the latter, we think that he dated the
 334 same unit as the dyke, which they possibly thought represented the Skalksepuit Granite
 335 intruding Draghoender Granite. The Steenkop Granite sample and the dyke sample plot close
 336 to each other in Figures 5 and 6, suggesting that they are related.

337

338 ***6.3 The Seekoebaard conglomerate.***

339 This sample DC1515 has a single age population comprising two zircon types, CL-dark and
 340 CL-bright. The concordia age of 2720 ± 4 Ma for these zircons shows that the conglomerate is
 341 not a xenolith in the older granites, but most likely a basal member of the Seekoebaard
 342 Formation, which has not been dated but is considered part of the Ventersdorp Supergroup.
 343

344 ***6.4 Geochemical Distinctions***

345 The validity of the normative distinctions of the two granite types was further tested.
 346 Diagrams involving Na, K, Ca and Sr all show the same grouping, the Draghoender type
 347 having lower potassium (0.7 - 1.8 compared with 3.6 – 4.9 K₂O wt%) and higher Na and Sr as
 348 shown in Figure 6. The two Steenkop-type samples are again clearly distinct with much lower
 349 Na₂O and low Sr.

350 The Maritzdam Granite sample DC1518, which is coeval with and classified as Skalksepups
 351 type, has much higher Sr content than the other samples in that group, suggesting a different
 352 magmatic evolution.

353

354 Rare Earth Element (REE) diagrams are shown in Figure 9. Both Draghoender and
 355 Skalksep groups show significant diversity but the groups are not clearly distinguished by
 356 REE geochemistry. The samples all show moderate to high Light REE enrichment and two

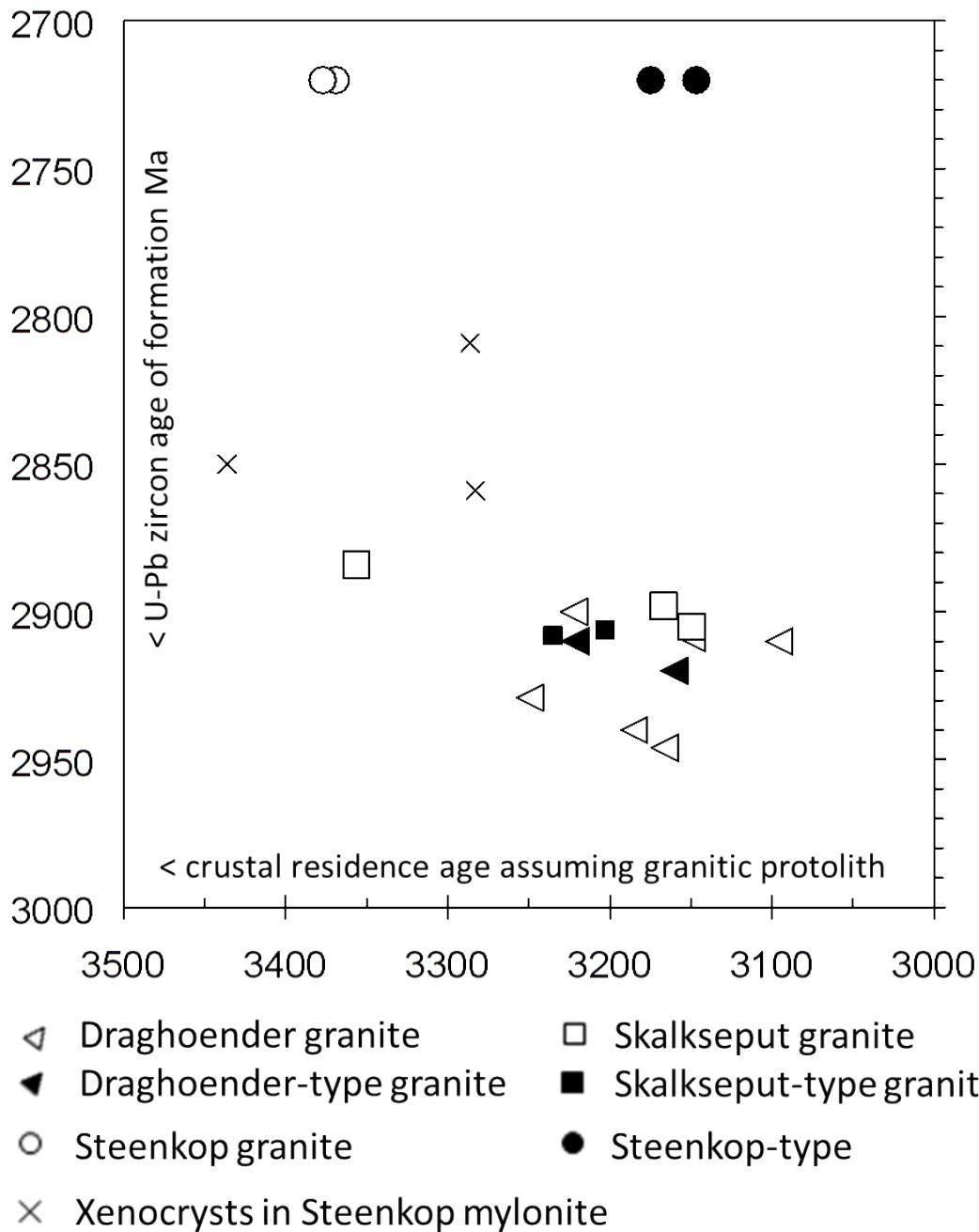
357 Draghoender type Trondhjemitic samples DC1332 and DC1613 show significant heavy REE
358 depletion to near-chondritic levels, which suggests that garnet was present in the (probably
359 mafic) protolith which melted to form these magmas. High-pressure, plagioclase-absent
360 melting processes for most samples are also suggested by the small to non-existent Eu
361 anomalies which both types show. The slope patterns of all the Skalksepups group samples are
362 similar, suggesting that they could be related by crystal fractionation which would generally
363 increase the concentrations without changing the slopes. By contrast the Draghoender group
364 samples have significantly different slope patterns and are unlikely to have been comagmatic,
365 which is also indicated by their 40 Ma range of zircon dates.

366

367 Lu-Hf data for zircons can be used to calculate crustal residence ages as shown in Figure 10,
368 which throw light on the crustal history of the area, although assumptions about the Depleted
369 Mantle source and Lu-Hf ratio of crustal protoliths are difficult to test. A granitic crustal
370 protolith with $^{176}\text{Lu}/^{177}\text{Hf}$ 0.015 gives the probable minimum crustal residence ages shown in
371 Figure 12 plotted against intrusion ages. Assuming basaltic crustal protolith with $^{176}\text{Lu}/^{177}\text{Hf}$
372 0.022 gives typically 200 Ma older crustal residence ages which may be valid for the
373 Draghoender type samples, thought to be generated by melting of eclogitic mafic crust in a
374 subduction setting. The Draghoender Figure 12 allows discussion of the origin of the samples
375 investigated. All but one of the Draghoender and Skalkseput (including type) samples have
376 similar crustal residence ages between 3.25 and 3.1 Ga, suggesting that they might have
377 similar crustal histories. The data also shows that the more evolved (potassic) Skalkseput
378 samples could have originated by melting of the plagioclase-rich Draghoender-like material
379 soon after it had formed. However as mentioned above, the tonalitic and trondhjemitic
380 Draghoender samples are more likely to have had mafic crustal protoliths which would
381 increase their crustal residence ages by about 0.2Ga, but not invalidate the suggestion above.
382 The exception is one Skalkseput sample DC01130 with crustal residence age at 3.35 Ga,
383 which suggests derivation from a different protolith.

384 The two Steenkop Granite samples have 0.2 Ga older crustal residence ages than the two
385 Steenkop-type samples which are the dyke DC1516 and the nearby conglomerate DC1515,
386 grouped with Steenkop because of their indistinguishable U-Pb ages and the similar
387 geochemistry of the dyke sample. One way in which their crustal histories can be reconciled
388 is if the Steenkop Granite samples were derived from a 3.4 Ga granitic protolith whereas the
389 Steenkop-type samples represent melted mafic crustal protolith of the same 3.4 Ga age. This

390 could be the case in view of the bimodal nature of the Ventersdorp Supergroup, with basaltic,
 391 andesitic and felsic units. Alternatively they were derived from protoliths which left the
 392 mantle at different times.



393

394

395 Figure 12. U-Pb zircon ages plotted against crustal residence ages assuming that the zircon
 396 originated from a magma which formed by melting of a granitic crustal protolith with
 397 $^{176}\text{Lu}/^{177}\text{Hf} = 0.015$.

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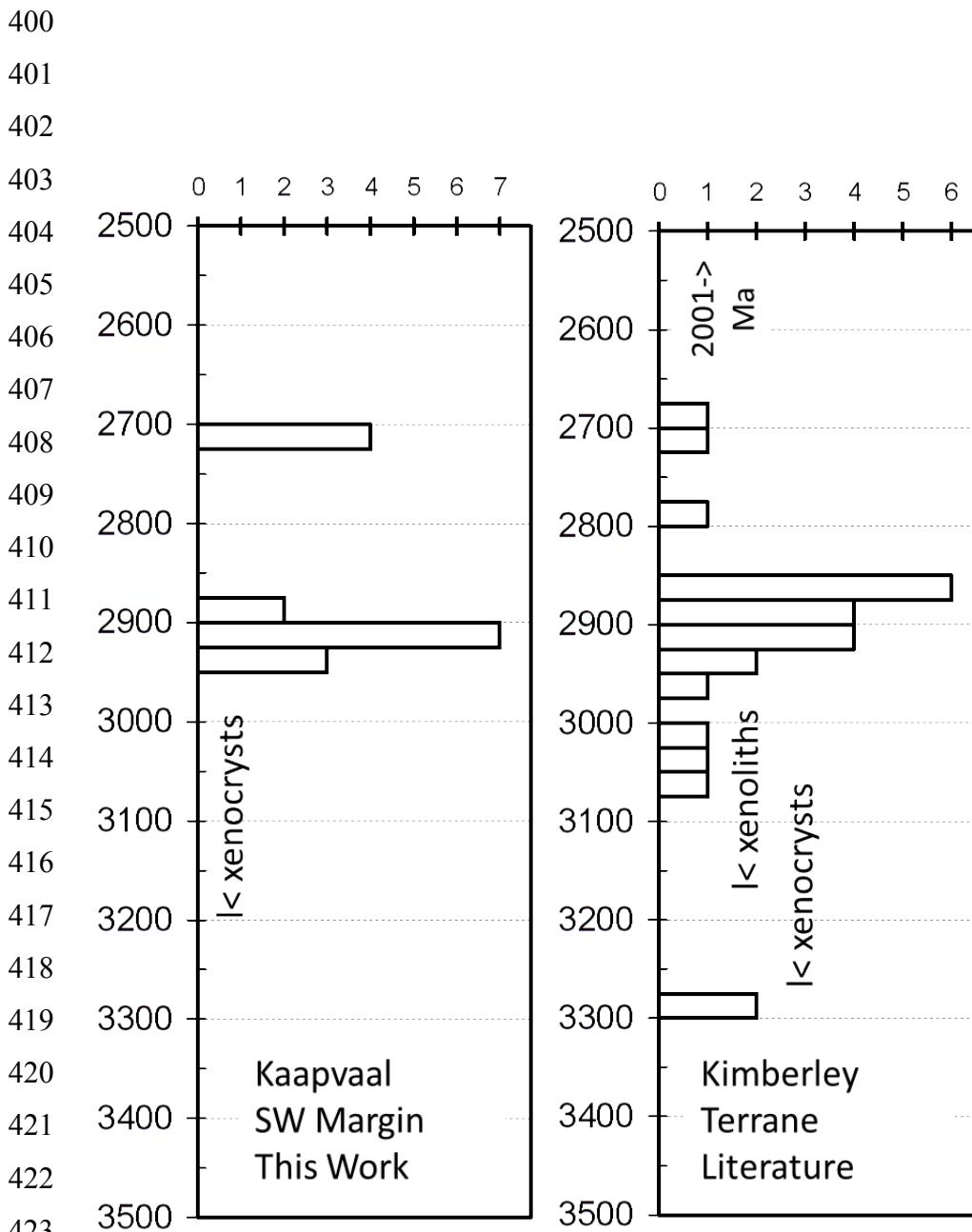


Figure 13. Comparison of precise ages for granitoids of the southwestern margin of the Kaapvaal Craton reported in this work with those from the literature for the Kimberley Terrane.

6.5 Regional implications.

As shown in Figure 13, the age distribution found in this work corresponds to a large extent with that for the Kimberley Terrane (equivalent to Block) of the Kaapvaal Craton extracted from the Dateview database, described by Eglington, and Armstrong (2004) and summarized in appendix A4. Granitic basement as old as 3280 Ma in the Kimberley mines and granitic

434 kimberlite xenoliths reflect the earliest well-documented crustal development in the terrane.
 435 Our smaller dataset yielded xenocrysts as old as 3180 Ma in the Draghoender Granite and
 436 even earlier crustal development is suggested by our Lu-Hf zircon crustal residence ages of
 437 3050 to 3450 Ma, shown in Figure 12. The main peak of emplacement ages between about
 438 2950 and 2850 Ma is thought to correspond to a subduction event followed by accretion of the
 439 Kimberley terrane to the greater Kaapvaal Craton (Schmitz et al., 2004). The ~2700 Ma event
 440 seen in both datasets can be ascribed to granitic magmatism related to the Ventersdorp
 441 volcanism.

442

443 ***6.6 Identity of the Marydale Terrane***

444 The rocks dated in this work clearly originated as part of the Archean Kimberley Terrane
 445 before becoming part of the greater Kaapvaal Craton. However their major uplift along the
 446 Doornberg Lineament, deformation and low-grade metamorphism, means that they are no
 447 longer part of the craton *sensu stricto* but should rather be regarded as part of the Marydale
 448 Terrane (Thomas et al., 1994, Hilliard, 1999). Weckmann et al., 2010 showed that high crustal
 449 resistivity characteristic of the Kaapvaal Craton continues far into the Namaqua-Natal
 450 Province in a profile southwest of Prieska. Their proposal that the craton boundary should be
 451 moved some 50 km south-westwards into the Namaqua-Natal Province may be interesting to
 452 diamond prospectors, but does not meet the requirement that a craton should be undeformed
 453 after its origin. It does seem possible that Kaapvaal cratonic basement could extend south-
 454 westwards beneath a major northeast-vergent thrust package which is exposed as the
 455 Marydale and Areachap terranes. Thus the cratonic limits may differ depending on the crustal
 456 level.

457

458 ***6.7 Significance of the Doornberg Fault***

459 All but one of our samples, the Maritzdam Granite, lie within the Doornberg Lineament *sensu*
 460 *lato* of Vajner (1974b), which corresponds to our Marydale Terrane between the Doornberg
 461 Fault *sensu stricto* and the Brakbos Shear Zone (Figure. 2). The Kaapvaal supracrustal cover
 462 sequence comprising the Ventersdorp and Griqualand West Supergroups do not occur
 463 southwest of the Doornberg Fault *sensu stricto*, which is why the main fault is shown in Fig. 2
 464 passing close to Marydale, contrary to Vajner (1974a) who shows it dying out into the
 465 Zeekoebaard volcanic rocks near Koegas. Our granite samples are from both sides of the fault
 466 and the question is whether they are all Kaapvaal basement granites. The Draghoender and
 467 Skalkseput Granite exposed in the Marydale High are clearly overlain by Kaapvaal

468 supracrustals and are Kaapvaal basement. The 2906 ± 8 Ma Maritzdam Granite (see Figures 1
 469 and 11) occupies the same stratigraphic position northeast of the Doornberg Lineament and is
 470 geochemically similar to and the same age as the 2902 ± 11 Ma Skalkseput granite samples,
 471 with which it should be correlated. The 2908 ± 4 Ma Welgevonden Granite sample south of
 472 the Doornberg Fault was too weathered to be geochemically analysed, but is also regarded as
 473 a Skalkseput correlate. Further north, the 2910 ± 8 Ma Franzenhof Granite corresponds
 474 geochemically to the Draghoender type and 2946 - 2900 Ma broad age range despite lying
 475 south of the Doornberg Fault. The third granite type exemplified by the younger 2719 ± 3 Ma
 476 Steenkop Granite, which crops out south of the Doornberg Fault is now known to have age
 477 correlates in the Marydale High in the form of 2720 ± 10 Ma dykes with similar
 478 geochemistry, as well as the conglomerate which contains only 2720 ± 4 Ma zircons of the
 479 same age and thought to represent the base of the Zeekoebaard Formation of the Ventersdorp
 480 Supergroup. Thus it seems that all three granite types straddle the Doornberg Fault which thus
 481 does not represent a major stratigraphic discontinuity.

482 Two questions remain. Why is the Kaapvaal supracrustal sequence absent south of the
 483 Doornberg Fault? Does the Marydale Group which (probably tectonically) overlies the
 484 Welgevonden and Franzenhof Granites represent a Ventersdorp correlate? These will be
 485 addressed in a later paper.

486

487 ***6.8 Crustal development of the Kimberley Terrane***

488 A possible scenario for the development of the Kimberley Terrane is as follows.

489 1. Crustal development began with mafic magmatism producing basaltic or komatiitic crust in
 490 a setting similar to a Phanerozoic mid-ocean ridge. Our zircon Lu-Hf crustal residence ages
 491 give clues to the timing of this event, which was before 3.2 Ga, assuming granitic crustal
 492 protoliths (Figure 12, or as early as 3.9 Ga, if a mafic protolith is assumed which seems likely
 493 for the trondhjemite rocks).

494 2. Granitic rocks formed as early as 3280 Ma in the Kimberley Mines area according to
 495 xenolith and xenocryst data from Kimberley (Dateview database described by Eglington, and
 496 Armstrong 2004). In the Marydale High felsic rocks containing zircon but otherwise of
 497 unknown composition formed as early as 3180 Ma evidenced by xenocrysts found in our
 498 samples.

499 3. Subduction of oceanic crust beneath the proto-Kimberley Terrane led to the formation of
 500 eclogites at depths around 30 km, melting of which gave rise to granitoid magmas which were
 501 emplaced between about 2950 Ma and 2890 Ma as plagioclase-rich trondjhemites and

502 tonalites of Draghoender type in the Marydale High area. This process gave rise to the bulk of
503 the granitic crust in the Kimberley Terrane.

504 4. Collision of the Kimberley Terrane with the Witwatersrand Terrane (Schmitz et al., 2004)
505 led to remelting of the tonalitic granitoids to produce K-rich granites of the Skalksepups type
506 at about 2915 – 2885 Ma, and resulted in the formation of a large craton of which the present
507 Kaapvaal Craton was a part.

508 5. A further thermal event originating in the mantle at about 2720 gave rise to huge basaltic
509 intrusions into the lower crust and after some crustal mixing, resulted in extrusions of
510 andesitic and felsic Ventersdorp lava over the entire Kaapvaal Craton, including the
511 Zeekoebaard lavas exposed in the Marydale High. Remelting of lower crustal granitoids
512 during this process also gave rise to granitic intrusions emplaced as the Steenkop type.

513

514 7. Conclusions

515 1. Three granitoid types, which crop out along the southwestern margin of the Kaapvaal
516 Craton, have been identified using geochemistry and microbeam zirconology.
517 2. The Draghoender type is tonalitic to trondhjemite in composition and ranges in age from
518 2946 ± 9 Ma to 2900 ± 6 Ma (n=6). It does not represent a single magma but is thought to have
519 been produced by melting of mafic crust at eclogite-facies depths during a subduction event.
520 3. The Skalksepups type (n=6) is granodioritic to quartz monzonitic, contains much more
521 (normative) alkali feldspar than the Draghoender type and has a weighted mean age of 2902 ± 11 Ma,
522 suggesting that the Skalksepups intrusions may represent a single magmatic event. It
523 is thought to correspond to the collision of the Kimberley Terrane with the Witwatersrand
524 Terrane.

525 4. The Steenkop Granite is monzogranitic and is about 200 Ma younger than the other types,
526 dated at 2719 ± 3 Ma. It lies athwart the Doornberg Fault which gave rise to mylonitic fabric
527 in one sample. It is geochemically similar to and coeval with 2720 ± 10 Ma dykes which
528 intrude the Skalksepup Granite in the Marydale High (Steenkop type n=3). A nearby
529 conglomerate with a single 2720 ± 4 Ma detrital zircon population is considered to be a basal
530 unit of the andesitic Zeekoebaard Formation. The Steenkop type is thus probably an intrusive
531 equivalent of the ~2720 Ma volcanic Ventersdorp Supergroup.

532 5. All but one of the samples studied lie within the 35 km-wide Marydale Terrane, bounded
533 by the Doornberg Fault and Brakbos Shear Zone, which is transitional between the
534 undeformed Kaapvaal Craton and the ~1.2 Ga tectonometamorphic Areachap Terrane of the
535 Namaqua-Natal Province. The granite samples we studied lie athwart the Doornberg Fault but

536 are all considered to have originated as part of the Kimberley Terrane basement which
 537 amalgamated with the Kaapvaal Craton at about 2900 Ma. The supracrustal rocks in the
 538 Marydale Terrane which overlie the granites are a northeast-vergent fold and thrust package,
 539 evidenced by their discontinuous stratigraphy and increasing metamorphic grade upwards in
 540 the stratigraphic sequence.

541

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 554 Geoscience, Pretoria, the rest were analysed by Riana Rossouw at the Central Analytical
 555 Facility, University of Stellenbosch.

556

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645 Captions for Figures

646 Figure 1. Precambrian structural provinces of southern Africa showing the locality of the
 647 study area (Figure. 2) at the southwestern margin of the Kaapvaal Craton. Selected precise U-
 648 Pb ages of other granitoids in the basement of the Kimberley block are shown.

649 Figure 2. Geological map showing the granites exposed on the southwestern margin of the
 650 Kaapvaal Craton. The Marydale Terrane lies between the Doornberg Fault and Brakbos Shear
 651 Zone. The Marydale High is the area underlain by the Draghoender and Skalkseput Granites
 652 between Marydale and Koegas. Sample localities south of the Marydale High are indicated as
 653 follows. Fr: Franzenhof Granite DC1517; St: Steenkop Granite Gneiss DC00121 from
 654 Steenkop; St my: Mylonitised Steenkop Granite DC01120 from Prieskapoort; Wel:
 655 Welgevonden Granite DC1333; Maritzdam Granite from the Kuip window of basement
 656 granite overlain by Ventersdorp Supergroup volcanic rocks.

657 Figure 3. Field pictures of samples taken in this work.

658 Figure 4. Precise U-Pb zircon dates for Kaapvaal basement granitoid samples exposed in the
 659 Marydale High, with localities marked by symbols Δ , \diamond and θ . Δ indicates Draghoender type
 660 tonalitic rocks with low QAP normative alkali feldspar (4-12%). \diamond indicates Skalkseput –
 661 type granodioritic rocks with higher QAP normative alkali feldspar (22-29%). θ indicates
 662 younger (2720 Ma) monzo-granite dykes coeval with the Steenkop Granite exposed near
 663 Prieskapoort. The sample numbers corresponding to ages can be found in Table x. Samples
 664 for which norms but not dates are available are shown with sample numbers. The H 2853 Ma
 665 age shown is by ion probe for Draghoender Granite from Hilliard 1999 and the A 2718 Ma
 666 age shown is an ion probe age by Richard Armstrong for his Skalksepups Granite, referred to
 667 in Hilliard 1999 and McCourt et al., 2000 and considered in this work to represent a
 668 Steenkop – aged dyke which we sampled at this locality. Precise locations for these two
 669 samples are not available. The dashed line is the Draghoender – Skalkseput boundary

according to the Council for Geoscience 1:250 000 Prieska map sheet of Malherbe and Moen 1991. Where it differs from that, the boundary according to Vajner 1974 is shown as a thin solid line.

Figure 5. Normative Quartz – Alkali feldspar – Plagioclase (QAP) mineral proportions for Granite samples from the southwestern margin of the Kaapvaal Craton including the Marydale high and outcrops southwards on both sides of the Doornberg Fault. Three types are distinguished as shown. The Draghoender and Skalksepup types overlap in U-Pb zircon age and the Steenkop type samples are significantly younger at ~2720 Ma.

Figure 6. Geochemical distinction of granite samples from the Southwestern Kaapvaal Craton. Three types are apparent, the Draghoender (triangles) and Skalksepups (blocks) granites which overlap in age and the younger ~2720 Ma Steenkop type.

Figure 7 Tectonic discrimination diagram for granites.

Fig. 8-1. Concordia diagrams showing the U-Pb zircon data and calculations for six samples from the Marydale High, arranged in order of latitude south, following Table x. Error ellipses are all shown at 1σ but age calculations at the 2σ confidence level

Fig. 8-2. Concordia diagrams showing the U-Pb zircon data and calculations for four samples from the Marydale High and two samples from further south, following Table x. Error ellipses are all shown at 1σ but age calculations at the 2σ confidence level.

Fig. 8-3. Concordia diagrams showing the U-Pb zircon data and calculations for four Granitoid samples and a conglomerate from the Marydale High. Error ellipses are all shown at 1σ .

Figure 9. Rare Earth Element (REE) diagrams for granitic samples, normalised to the chondritic values of Evensen (1978).

Figure 10. Zircon Lu-Hf data for 12 samples plotted on a Hf isotope evolution diagram. The calculation of crustal residence ages for a sample is illustrated by projecting an evolution line corresponding to a hypothetical granitic precursor with $^{176}\text{Lu}/^{177}\text{Hf}$ 0.015 from the data point back to intersect with the Depleted Mantle line, in this case at 3.5 Ga. More mafic precursors with e.g. $^{176}\text{Lu}/^{177}\text{Hf}$ 0.022 would give older crustal residence ages.

Figure 11. Precise U-Pb zircon microbeam dates for granitic samples from the southwest margin of the Kaapvaal Craton arranged in order of latitude south. Sample numbers are followed by Taken As - Classified As. e.g D – S means taken as Draghoender, classified as Skalksepup type. St means Steenkop type.

Figure 12. U-Pb zircon ages plotted against crustal residence ages assuming that the zircon originated from a magma which formed by melting of a granitic crustal protolith with $^{176}\text{Lu}/^{177}\text{Hf}$ 0.015.

Figure 13. comparison of precise ages for granitoids of the southwestern margin of the Kaapvaal Craton reported in this work with those from the literature for the Kimberley Terrane.

708

709 Captions for Tables

710 Table 1. Major element analyses and normative data for granitoid samples from the
711 Southwestern margin of the Kaapvaal Craton.

712 Table 2. Summary of U-Pb zircon dating results for granitoid from the southwestern margin
713 of the Kaapvaal Craton. Methods abbreviations are: LA Laser Ablation; HR High Resolution;
714 MC multicollector; ICPMS Inductively coupled Mass Spectrometry; SIMS secondary ion
715 mass spectrometry.

716

717 **Ancillary Data appendices**

718 A1 Full geochemical data.

719 A2 Microbeam U-Pb zircon data.

720 A3 Lu-Hf zircon data.

721 A4 U-Pb zircon age data for the Kimberley Terrane of the Kaapvaal Craton extracted from the
722 Dateview database, described by Eglington, and Armstrong (2004).

723

U-Pb microbeam dating

Sample	Unit	method
DC01120	Steenkop Granite Gneiss	discordia SIMS
DC01120	Steenkop Granite Gneiss	discordia LA MCICPMS
DC01121	Steenkop Granite Mylonite	Pb-Pb age LA MCICPMS
DC01130	Skalkseput Granite	discordia LA MCICPMS
DC1332	Draghoender Granite	discordia SIMS
DC1333	Welgevonden Granite	discordia SIMS
DC1510	Draghoender Granite	discordia LA HR ICPMS
DC1512	Skalkseput Granite	discordia LA HR ICPMS
DC1513	Skalkseput Granite	discordia LA HR ICPMS
DC1515	conglomerate Zeekoebaard	discordia LA HR ICPMS
DC1516	Dyke in Skalkseput	discordia LA HR ICPMS
DC1517	Franzenhof Granite	discordia LA HR ICPMS
DC1518	Ongers River Granite	concordia LA HR ICPMS
DC1610	Draghoender Granite	discordia LA HR ICPMS
DC1611	Draghoender Granite	discordia LA HR ICPMS
DC1612	Skalkseput Granite	discordia LA HR ICPMS
DC1613	Skalkseput Granite	discordia LA HR ICPMS classified as Draghoender type

DC01120	Lat.S deg	S Min	Lon.E deg	E Min															
Steenkop Granite Gneiss	29	48.628	22	41.26															
	Laser Ablation Multicollector ICPMS University of Oslo		Puck:	DHC0301															
	ppm	Ratios			Discordance														
Grain	Age	Domain	U	^{206}Pb	$206/204\text{Pb}$	$^{207}\text{Pb}/^{206}\text{Pb}$	1SE	$^{207}\text{Pb}/^{235}\text{U}^*$	1SE	$^{206}\text{Pb}/^{238}\text{U}^*$	1SE	Rho	Central (%)	Minimum rim (%)	Ages				
1 Main	16		8.3	1014	0.1886	0.00088	13.04092	0.30956	0.5015	0.011673	0.981	-4.9	-3.6	2730	7	2683	22	2620	50
2 xenocryst	83		42.1	4188	0.20045	0.00075	13.27152	0.1412	0.480188	0.00478	0.936	-12.9	-11.9	2830	6	2699	10	2528	21
4 Main	284		157.4	13255	0.18694	0.00046	13.54682	0.12516	0.525583	0.004681	0.964	0.3 .		2715	4	2719	9	2723	20
6 Main	75		41.2	10370	0.18748	0.00054	13.63432	0.14124	0.527446	0.005246	0.960	0.5 .		2720	5	2725	10	2731	22
8 Main*	242		115.9	7263	0.18874	0.0005	11.89643	0.11425	0.457144	0.004221	0.961	-13.3	-12.7	2731	4	2596	9	2427	19
9 Main*	7		3.4	713	0.18708	0.00194	13.20906	1.18136	0.511762	0.045491	0.993	-2.4 .		2717	16	2694	84	2664	194
10 Main	29		14.1	1509	0.18512	0.0008	12.21646	0.17937	0.47863	0.006716	0.956	-8	-6.7	2699	7	2621	14	2521	29
15 Main	218		110.9	10306	0.18752	0.00048	12.57483	0.13832	0.486352	0.005203	0.972	-7.4	-6.7	2721	4	2648	10	2555	23
16 Main	40		19.4	1822	0.18671	0.00086	12.4908	0.20153	0.485191	0.007506	0.959	-7.3	-6	2713	7	2642	15	2550	33

* excluded from regression

const,err.

DC01130		Lat.S deg	S Min	Lon.E deg	E Min	2012/12/01														
Skalkseput Granite		29	48.628	22	41.26															
Laser Ablation Multicollector ICPMS University of Oslo Analysed by Magnus Kristoffersen																				
Grain	Age Domain	ppm	U	206Pb	206/204	207Pb/206Pb*	1SE	207Pb/235U*	1SE	206Pb/238U*	1SE	Rho	Central (%)	Minimum rim (%)	Discordance	Ages				
															207/206	1σ	207/235	1σ	206/238	1σ
44	main		677	148.1	1250	0.18214	0.00215	4.86022	0.12054	0.193529	0.004223	0.88	-62.3	-61	2672	19	1795	21	1140	23
13*	main		1206	256.5	3842	0.16905	0.00195	4.37415	0.11821	0.187658	0.004586	0.904	-61.3	-59.9	2548	19	1707	22	1109	25
26	main		523	127.2	1218	0.18805	0.00228	5.49432	0.14449	0.2119	0.004943	0.887	-59.7	-58.3	2725	19	1900	23	1239	26
42*	main		1066	266.5	794	0.19069	0.00232	5.78155	0.13909	0.219898	0.004564	0.863	-58.6	-57.2	2748	19	1944	21	1281	24
23	main		550	167.5	1339	0.19408	0.00237	7.11145	0.21026	0.265755	0.007155	0.911	-50.6	-48.8	2777	20	2125	26	1519	36
7	main		777	236.8	2192	0.19068	0.0023	7.02055	0.1994	0.267029	0.006868	0.906	-49.8	-47.9	2748	19	2114	25	1526	35
4	main		611	207.1	1060	0.19397	0.00236	7.85173	0.19887	0.293588	0.006519	0.877	-45.5	-43.5	2776	19	2214	23	1659	32
25	main		329	129.5	2431	0.19695	0.00244	9.19829	0.25777	0.338735	0.008514	0.897	-37.8	-35.5	2801	19	2358	26	1881	41
27	main		546	215	11603	0.19444	0.00237	9.1	0.25774	0.339274	0.008678	0.903	-37.1	-34.8	2780	19	2348	26	1883	42
30	main		399	161	20912	0.19514	0.00239	9.31864	0.2655	0.346343	0.008913	0.903	-35.9	-33.6	2786	19	2370	26	1917	43
10	main		317	134.1	1025	0.20169	0.00253	10.01809	0.31855	0.360238	0.010525	0.919	-34.9	-32.5	2840	19	2436	29	1983	50
2	main		291	131.5	2621	0.20359	0.00259	10.78128	0.32183	0.384071	0.010375	0.905	-31.1	-28.5	2855	20	2504	28	2095	48
3	main		256	115.2	3178	0.20349	0.00257	10.87917	0.34923	0.387756	0.01144	0.919	-30.4	-27.8	2854	20	2513	30	2112	53
40	main		90	46.2	1168	0.20935	0.00271	12.59648	0.39699	0.436394	0.012536	0.911	-23.2	-20.3	2901	20	2650	30	2334	56
49	main		71	41.5	2069	0.20321	0.00265	13.84471	0.47761	0.494122	0.015781	0.926	-11.2	-7.7	2852	21	2739	33	2588	68

* excluded from regression

DC1610 Draghoender Granite

Lat. S deg 30 S Min 8.5518 E deg 23 E Min 13.503

Laser Ablation high resolution ICPMS

University of Stellenbosch

RATIOS**AGES [Ma]****Conc.**

grain	Age domain	U [ppm] ^a	Pb [ppm] ^a	Th [ppm]	$^{206}\text{Pb}/^{204}\text{Pb}$	Th/U meas	$^{207}\text{Pb}/^{235}\text{U}$ ^b	$1\sigma^d$	$^{206}\text{Pb}/^{238}\text{U}$ ^b	$1\sigma^d$	ρho^c	$^{207}\text{Pb}/^{206}\text{Pb}$ ^e	$1\sigma^d$	$^{207}\text{Pb}/^{235}\text{U}$	1σ	$^{206}\text{Pb}/^{238}\text{U}$	1σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1σ	%
3	main osc zoned	137	72	48	593	0.35	15.64	0.25	0.5293	0.0061	0.71	0.21427	0.00246	2855	46	2739	31	2938	19	93
7.1	main osc zoned	298	146	138	1613	0.46	14.50	0.23	0.4880	0.0056	0.71	0.21548	0.00244	2783	45	2562	29	2947	18	87
7.2	main osc zoned	335	120	72	1603	0.22	10.59	0.17	0.3593	0.0042	0.70	0.21385	0.00252	2488	41	1979	23	2935	19	67
8	main osc zoned	217	120	90	2112	0.42	16.41	0.26	0.5505	0.0063	0.71	0.21619	0.00246	2901	47	2827	32	2953	18	96
10	main osc zoned	235	136	92	2133	0.39	17.16	0.28	0.5774	0.0066	0.71	0.21559	0.00245	2944	47	2938	34	2948	18	100
11	main osc zoned	554	315	145	3544	0.26	16.73	0.27	0.5680	0.0065	0.71	0.21367	0.00241	2920	47	2899	33	2934	18	99
13	main osc zoned	188	108	52	1952	0.28	17.24	0.28	0.5764	0.0066	0.71	0.21693	0.00249	2948	48	2934	34	2958	19	99
16	main osc zoned	299	139	92	3383	0.31	13.80	0.25	0.4637	0.0056	0.66	0.21581	0.00300	2736	50	2456	30	2950	22	83
19	main osc zoned	152	77	34	2173	0.22	15.19	0.26	0.5095	0.0060	0.68	0.21627	0.00272	2827	49	2655	31	2953	20	90
20	main osc zoned	239	121	95	1774	0.40	14.93	0.24	0.5042	0.0058	0.71	0.21474	0.00248	2811	46	2632	30	2942	19	89
22	main osc zoned	214	123	96	2478	0.45	17.16	0.28	0.5741	0.0066	0.70	0.21682	0.00253	2944	48	2925	34	2957	19	99
26	main osc zoned	270	156	79	2395	0.29	17.23	0.28	0.5791	0.0067	0.70	0.21582	0.00251	2948	48	2945	34	2950	19	100
28	main osc zoned	230	133	101	3213	0.44	17.18	0.28	0.5788	0.0067	0.70	0.21529	0.00254	2945	49	2944	34	2946	19	100
48	main osc zoned	330	164	123	2536	0.37	14.82	0.25	0.4969	0.0058	0.69	0.21631	0.00261	2804	47	2601	30	2953	19	88
49	main osc zoned	266	124	95	2234	0.36	13.76	0.24	0.4643	0.0055	0.68	0.21487	0.00272	2733	47	2459	29	2943	20	84
50	main osc zoned	149	86	60	2213	0.40	17.11	0.30	0.5794	0.0068	0.68	0.21418	0.00275	2941	51	2946	35	2937	21	100
54	xenocryst	119	52	41	519	0.34	13.72	0.26	0.4381	0.0054	0.66	0.22708	0.00322	2730	51	2342	29	3032	23	77
57	main osc zoned	364	202	146	3670	0.40	16.53	0.28	0.5557	0.0064	0.69	0.21572	0.00261	2908	49	2849	33	2949	20	97
61	main osc zoned	131	73	26	869	0.20	16.34	0.30	0.5548	0.0065	0.63	0.21359	0.00304	2897	53	2845	33	2933	23	97
66	main osc zoned	114	66	46	1670	0.40	17.01	0.30	0.5783	0.0068	0.66	0.21331	0.00283	2935	52	2942	35	2931	21	100
69	main osc zoned	210	121	77	2733	0.37	17.18	0.29	0.5763	0.0067	0.68	0.21624	0.00270	2945	50	2933	34	2953	20	99
74.1	main osc zoned	282	163	130	3911	0.46	17.11	0.29	0.5771	0.0067	0.68	0.21497	0.00268	2941	50	2937	34	2943	20	100
74.2	main osc zoned	363	210	212	8050	0.59	17.06	0.29	0.5780	0.0067	0.68	0.21406	0.00268	2938	50	2941	34	2937	20	100
75	main osc zoned	220	127	79	3287	0.36	17.20	0.30	0.5754	0.0067	0.67	0.21679	0.00278	2946	51	2930	34	2957	21	99
79	main osc zoned	252	145	140	5055	0.56	17.12	0.30	0.5766	0.0067	0.67	0.21531	0.00275	2941	51	2935	34	2946	21	100
82	main osc zoned	80	43	16	1459	0.20	15.66	0.29	0.5330	0.0064	0.65	0.21308	0.00299	2856	53	2754	33	2929	23	94
83.1	main osc zoned	293	105	52	598	0.18	10.77	0.20	0.3582	0.0043	0.64	0.21803	0.00316	2503	47	1973	24	2966	23	67
83.2	main osc zoned	617	211	120	2897	0.19	10.08	0.18	0.3420	0.0040	0.66	0.21369	0.00284	2442	43	1896	22	2934	21	65
83.3	main osc zoned	345	184	46	5254	0.13	15.79	0.28	0.5323	0.0062	0.66	0.21509	0.00289	2864	51	2751	32	2944	22	93

^aU and Pb concentrations and Th/U ratios are calculated relative to GJ-1 reference zircon^bCorrected for background and within-run Pb/U fractionation and normalised to reference zircon GJ-1 (ID-TIMS values/measured value); ^c $^{207}\text{Pb}/^{235}\text{U}$ calculated using $(^{207}\text{Pb}/^{206}\text{Pb})/(^{238}\text{U}/^{206}\text{Pb} * 1/137.88)$ ^dRho is the error correlation defined as the quotient of the propagated errors of the $^{206}\text{Pb}/^{238}\text{U}$ and the $^{207}\text{Pb}/^{235}\text{U}$ ratio^eQuadratic addition of within-run errors (2 SD) and daily reproducibility of GJ-1 (2 SD)^fCorrected for mass-bias by normalising to GJ-1 reference zircon (~0.6 per atomic mass unit) and common Pb using the model Pb composition of Stacey & Kramers (1975)

Highlights

Three granitoid types identified in SW Kaapvaal Craton basement.

2946-2900 Ma tonalite-trondhjemite Draghoender type are mafic subduction-melts.

2902 ±11 Ma Skalkseput type reflect Kimberley - Witwatersrand terrane collision.

2719 ±3 Ma Ventersdorp-age Steenkop monzogranite lies athwart the Doornberg Fault.

Kaapvaal basement continues SW of Doornberg Fault beneath Namaqua foreland thrusts.